

**United States Department of Agriculture  
Soil Conservation Service**

**Chester, PA 19013**

**Subject:** Ground-Penetrating Radar (GPR)  
study of depth to paleosols in Reno  
County, Kansas; 20 to 30 June 1994.

**Date:** 28 July 1994

**To:** Richard L. Schlepp  
State Soil Scientist  
USDA - Soil Conservation Service  
760 S. Broadway  
Salina, Kansas 67401

**Purpose:**

To use GPR and EM techniques to characterize depths to water table, lamellae and paleosols, and salt contents of selected soils and map units in Reno County, Kansas.

**Participants:**

- Brian Bergentine, Civil Engineer, SCS, Hutchinson, KS
- Andy Broxterman, Agricultural Engineer, SCS, Hutchinson, KS
- Ricky Cox, Area Resource Soil Scientist, SCS, Hutchinson, KS
- Steve Graber, Area Soil Scientist, SCS, Dodge City, KS
- Jim Doolittle, Soil Specialist, SCS, Chester, PA
- Mary Doolittle, Earth Team Volunteer, NNTC, Chester, PA
- Paul Finnell, Ass't. State Soil Scientist, SCS, Salina, KS
- David Kacirik, Soil Scientist, SCS, Leavenworth, KS
- Jim McDowell, Soil Survey Party Leader, SCS, Liberal, KS
- Bob Murphy, Soil Survey Party Leader, SCS, Hutchinson, KS
- Gary Parks, Soil Scientist, SCS, Hutchinson, KS
- Mike Sucik, Soil Survey Party Leader, SCS, Leavenworth, KS
- Robert Triggs, Area Resource Soil Scientist, SCS, Hays, KS
- John Warren, Soil Scientist SCS, Hutchinson, KS
- Bill Wehmuller, Area Resource Soil Scientist, SCS, Manhattan, KS

**Activities:**

During this field study, about 24 miles of continuous radar data were collected. Traverses were completed with GPR in areas of Carwile (fine, mixed, thermic Typic Argiaquoll), Dillwyn (mixed, thermic Aquic Ustipsamment), Pratt (sandy, mixed, thermic Psammentic Haplustalf), and Tivoli (mixed, thermic Typic Ustipsamment) soils. Traverses were also conducted in areas of the proposed Dillhut (sandy over loamy, mixed, thermic Aquic Ustorthent), Hayes (coarse-loamy, mixed, thermic Udic Haplustalf) Langdon (mixed, thermic Argic Ustipsamment), Plev (sandy, mixed, thermic Mollic Endoaqualf), and Solvay (fine-loamy, mixed, thermic Aquic Haplustalf) soils. These traverses were conducted principally in the northeast and west portions of Reno County.

A detailed GPR and EM site investigation was completed in an area of Langdon-Tivoli fine sands, 0 to 15 percent slopes. Several transects and a detailed grid were completed with the EM38 meter in areas of Tabler (fine, montmorillonitic, thermic Udertic Argiustoll) and the proposed Punkin (fine, mixed, thermic Aquic Natrustoll) soils.

**Equipment:**

The radar unit used in this study was the Subsurface Interface Radar (SIR) System-8 manufactured by Geophysical Survey Systems, Inc<sup>+</sup>. The system was powered by a 12-volt vehicular battery. The model 3110 (120 MHz) antenna with model 705DA transceiver was used in the investigation.

The electromagnetic induction meter was the EM38, manufactured by Geonics Limited<sup>+</sup>. The meter is portable and requires only one person to operate. The depth of penetration is dependent upon the intercoil spacing, transmission frequency, and coil orientation relative to the ground surface. The EM38 meter integrates values of apparent conductivity over the upper 0.75 m in the horizontal dipole orientation, and over the upper 1.5 m in the vertical dipole orientation.

**Discussion:****Ground-Penetrating Radar Transects:****Field Procedures:**

All transects were conducted in areas that had been preselected by the survey party. Transects were confined to delineated areas of recognized soil map units. A 100 foot interval (paced) was used between observation points. Though GPR provides a continuous record of the subsurface, estimates of the depths to paleosols were restricted to the observation points.

Calibration trials consisted of multiple traverses conducted across site 1A (see Table 1). A scanning time of 80 nanoseconds (ns) was used. As part of the calibration trials, a metallic reflector was buried at a depth of 17 inches (43 cm). Based on the scaled depth to this reflector, the calculated dielectric constant was 9.9. The velocity of propagation was 0.312 ft/ns (0.095 m/ns). With a scanning time of 80 ns, the observation depth was about 12.5 feet (3.81 m).

The depth of observation ranged from about 17 to 150 inches. Because of the strong reflected signal from the soil surface, features within a depth of 17 inches of the soil surface were poorly or not resolved with the 120 MHz antenna. Because of the restricted scanning time, features at depths greater than 150 inches were beyond the maximum observation depth.

During the course of this study, at twenty-six observation sites, a soil auger was used to determine the depth to paleosol. This information was used to verify interpretations and to scale the radar profiles. The correlation between auger observations and scaled radar depths was exceptionally high. Based on 26 observations, the coefficient of determination ( $r^2$ ) between the observed auger and the interpreted radar depths to the paleosol was 0.9593 (see Figure 1). The average difference in the depth to paleosol measurements between soil auger and radar measurements was 2.62 inches (with a range of 0 to 8 inches).

---

<sup>+</sup> Trade names have been used to provide specific information. Their mention does not constitute endorsement.

Ground-penetrating radar techniques are not appropriate for soil investigations in many areas of Reno County. This study was restricted to map units with components of coarse and moderately-coarse textured soils. However, in these units, depths of penetration were often restricted by thick bands of lamellae or layers of coarse-loamy or finer textured soil materials. In most areas of coarse-textured soil materials, GPR was a suitable investigation and research tool and provided adequate depths of penetration and good resolution of subsurface interfaces. However, in areas of coarse-textured materials, while the depth of observation and the resolution of subsurface features were good, the number and segmented nature of the subsurface layers plagued interpretations. To properly identify each layer, a prohibitive number of auger observations would have been required. One layer, the "paleosol", was distinct and recognizable in many areas and across most radar profiles.

The "paleosol" was defined as the lowest layer which limited the observation depth of the radar. Generally, this layer consisted of sandy loam or finer textured materials. Below this layer, no imagery was observed because of the high rates of signal attenuation. The paleosol appeared continuous on most radar profiles. However, in some areas, this layer was discontinuous. In addition, auger observations taken along the same transect line, confirmed that the texture of this layer can grade laterally.

The radar profiles contained numerous, segmented layers above the paleosols. These layers were assumed to represent stratified sand layers, buried A horizons, lamellae, or the water table. However, as ground-truth auger observations were limited, the identity of most of these layers could not be confirmed.

The map units, transect lengths, and locations are listed in Table 1. Ground-penetrating radar data were collected along eighty-five transect lines which ranged in length from 100 to 5400 feet. Several transects were omitted from this study because of limited profiling depths or the poor resolution of subsurface interfaces.

#### Results:

Transects were completed with the following map units: Ax - Hayes-Solvay complex, 0 to 5 percent slopes; Dp - Dillhut-Plev complex, 0 to 2 percent slopes; Ds - Dillhut-Solvay complex, 0 to 3 percent slopes; Dt - Dillwyn-Tivoli fine sands, 0 to 15 percent slopes; Pr - Pratt loamy fine sand, 1 to 5 percent slopes; Pt - Turon-Solvay-Carwile complex, 0 to 5 percent slopes; Px - Langdon-Tivoli fine sands, 0 to 15 percent slopes; and Tf - Tivoli fine sand 10 to 30 percent slopes. Table 2 summarizes the transect data for each map unit. Data from each transect are listed in the Addendum to this report.

Table 2 conveys comparative information but the use of averages and ranges is statistically inappropriate. The use of these statistics is incorrect as the minimum and maximum observation depths were 17 and 150 inches, respectively. Regardless of actual depths, because of the limitations of the equipment and the constraints of the survey, paleosols occurring at depths of less than 17 inches or greater than 150 inches were recorded as 17 and 150 inches, respectively.

In areas of the Hayes-Solvay complex, 0 to 5 percent slopes (map unit Ax), the average depth to the paleosol was 38.7 inches. One-half the observations had depths to paleosol between 33 and 42 inches. In the transected area of this map unit, depths to the paleosol were shallow

**Table 1**  
**LOCATION OF GPR TRANSECTS**

<b>Number</b>	<b>Symbol</b>	<b>Length (feet)</b>	<b>Location</b>
1A	DT	900	SW1/4 SEC 11, T. 22 S., R. 6 W.
1	DP	1700	SW1/4 SEC 11, T. 22 S., R. 6 W.
2	DT	1100	S 1/2 SEC 11, T. 22 S., R. 6 W.
3	DT	1200	S 1/2 SEC 11, T. 22 S., R. 6 W.
4	TF	300	SE1/4 SEC 11, T. 22 S., R. 6 W.
5	DT	600	SE1/4 SEC 11, T. 22 S., R. 6 W.
6	DP	600	SE1/4 SEC 11, T. 22 S., R. 6 W.
7	DP	2400	SE1/4 SEC 5, T. 22 S., R. 5 W.
8	PR	100	SE1/4 SEC 5, T. 22 S., R. 5 W.
9	DP	2600	SE1/4 SEC 5, T. 22 S., R. 5 W.
10	DT	2500	NW1/4 SEC 5, T. 22 S., R. 5 W.
11	DS	2600	SW1/4 SEC 10, T. 22 S., R. 6 W.
12	PR	2500	S 1/2 SEC 15, T. 22 S., R. 5 W.
13	DT	2600	NW1/4 SEC 14, T. 22 S., R. 5 W.
14	DT	2100	SW1/4 SEC 13, T. 22 S., R. 5 W.
15	DT	2800	S 1/2 SEC 13, T. 22 S., R. 5 W.
16	PR	800	NW1/4 SEC 19, T. 22 S., R. 5 W.
17	DT	1900	NW1/4 SEC 19, T. 22 S., R. 5 W.
18	DT	3100	NE1/4 SEC 19, T. 22 S., R. 5 W.
19	TF	200	SW1/4 SEC 21, T. 22 S., R. 5 W.
20	DT	1000	SW1/4 SEC 21, T. 22 S., R. 5 W.
21	DP	1200	SW1/4 SEC 21, T. 22 S., R. 5 W.
22	DT	700	SW1/4 SEC 21, T. 22 S., R. 5 W.
23	DP	400	NE1/4 SEC 21, T. 22 S., R. 5 W.
24	TF	400	NE1/4 SEC 21, T. 22 S., R. 5 W.
25	DT	300	SW1/4 SEC 22, T. 22 S., R. 5 W.
26	DS	1000	SE1/4 SEC 22, T. 22 S., R. 5 W.
27	DT	900	SE1/4 SEC 22, T. 22 S., R. 5 W.
28	TF	300	SE1/4 SEC 22, T. 22 S., R. 5 W.
29	DS	200	SE1/4 SEC 22, T. 22 S., R. 5 W.
30	TF	100	SE1/4 SEC 22, T. 22 S., R. 5 W.
31	DS	100	SE1/4 SEC 22, T. 22 S., R. 5 W.
32	DP	1300	SW1/4 SEC 19, T. 22 S., R. 4 W.
33	PR	300	SW1/4 SEC 19, T. 22 S., R. 4 W.
34	DS	400	SW1/4 SEC 19, T. 22 S., R. 4 W.
35	DS	1100	SE1/4 SEC 24, T. 22 S., R. 5 W.
36	DT	1200	SE1/4 SEC 24, T. 22 S., R. 5 W.
37	DS*	1600	SW1/4 SEC 24, T. 22 S., R. 5 W.
38	DS	1200	SW1/4 SEC 30, T. 22 S., R. 4 W.
39	PT	2000	SE1/4 SEC 30, T. 22 S., R. 4 W.
40	DS	1100	SW1/4 SEC 28, T. 22 S., R. 4 W.
41	DS	1200	NW1/4 SEC 1, T. 23 S., R. 4 W.

\* May contain more than one map unit.

Table 1 (continued)

## LOCATION OF GPR TRANSECTS

Number	Symbol	Length (feet)	Location				
42	DT	1100	NW1/4	SEC 1,	T. 23	S., R. 4 W.	
43	DS	1700	NE1/4	SEC 1,	T. 23	S., R. 4 W.	
44	DS	1500	NE1/4	SEC 11,	T. 23	S., R. 4 W.	
45	DT	4800	S 1/2	SEC 11,	T. 23	S., R. 4 W.	
46	DS	1500	NE1/4	SEC 14,	T. 23	S., R. 4 W.	
47	DS	2300	NE1/4	SEC 13,	T. 23	S., R. 4 W.	
48	DT	3600	N 1/2	SEC 7,	T. 23	S., R. 4 W.	
49	DT	800	SW1/4	SEC 1,	T. 23	S., R. 5 W.	
50	DS	900	SW1/4	SEC 1,	T. 23	S., R. 5 W.	
51	DT	1600	SW1/4	SEC 1,	T. 23	S., R. 5 W.	
52	TF	1400	SW1/4	SEC 1,	T. 23	S., R. 5 W.	
53	DS	1000	NW1/4	SEC 4,	T. 23	S., R. 4 W.	
54	DT	1300	NE1/4	SEC 4,	T. 23	S., R. 4 W.	
55	DT	1500	SW1/4	SEC 9,	T. 23	S., R. 4 W.	
56	TF	800	SW1/4	SEC 9,	T. 23	S., R. 4 W.	
57	DT	1100	SW1/4	SEC 9,	T. 23	S., R. 4 W.	
58	DS	2700	NE1/4	SEC 17,	T. 23	S., R. 4 W.	
59	DS	1700	SW1/4	SEC 17,	T. 23	S., R. 4 W.	
60	DS	3400	SE1/4	SEC 18,	T. 23	S., R. 4 W.	
61	DS*	3900	S 1/2	SEC 20,	T. 22	S., R. 10 W.	
62	DT	5400	S 1/2	SEC 21,	T. 22	S., R. 10 W.	
63	DT	1600	SW1/4	SEC 8,	T. 22	S., R. 10 W.	
64	DS	1200	SE1/4	SEC 8,	T. 22	S., R. 10 W.	
65	TF	1500	SE1/4	SEC 8,	T. 22	S., R. 10 W.	
66	DS	1400	SW1/4	SEC 5,	T. 22	S., R. 10 W.	
67	DT	500	SW1/4	SEC 5,	T. 22	S., R. 10 W.	
68	DT	1400	NE1/4	SEC 5,	T. 22	S., R. 10 W.	
69	TF	500	NE1/4	SEC 5,	T. 22	S., R. 10 W.	
70	DS	1000	NE1/4	SEC 5,	T. 22	S., R. 10 W.	
71	DT	800	NE1/4	SEC 5,	T. 22	S., R. 10 W.	
72	DS	300	NE1/4	SEC 5,	T. 22	S., R. 10 W.	
73	DT	1300	NE1/4	SEC 5,	T. 22	S., R. 10 W.	
74	PX	3900	N 1/2	SEC 8,	T. 24	S., R. 10 W.	
75	PR	500	NW1/4	SEC 8,	T. 24	S., R. 10 W.	
76	PX	300	NE1/4	SEC 3,	T. 24	S., R. 9 W.	
77	PT	1100	NE1/4	SEC 3,	T. 24	S., R. 9 W.	
78	PX	600	NE1/4	SEC 3,	T. 24	S., R. 9 W.	
79	PT	1200	N 1/2	SEC 3,	T. 24	S., R. 9 W.	
80	AX	1400	NW1/4	SEC 3,	T. 24	S., R. 9 W.	
81	PX	2400	NW1/4	SEC 11,	T. 24	S., R. 9 W.	
82	PX	2100	N 1/2	SEC 7,	T. 25	S., R. 9 W.	
83	PX	1900	NE1/4	SEC 13,	T. 25	S., R. 9 W.	
84	PX	3400	N 1/2	SEC 13,	T. 25	S., R. 10 W.	
85	PX	1200	SW1/4	SEC 10,	T. 25	S., R. 10 W.	

at 7 percent, moderately deep at 67 percent, deep at 20 percent, and very deep at 6 percent of the observation points (see Figure 2, upper).

In areas of the Dillhut-Plev complex, 0 to 2 percent slopes (map unit Dp), the average depth to the paleosol was 28.9 inches. One-half the observations had depths to paleosol between 17 and 34 inches. In the transected area of this map unit, depths to the paleosol were shallow at 52 percent, moderately deep at 30 percent, deep at 11 percent, and very deep at 7 percent of the observation points (see Figure 2, lower).

**Table 2**  
**GPR TRANSECTS**

Map Unit	Number	Observations	Length(ft)	Depth to Paleosol (in inches)		
				Min.	Max.	Avg.
Ax	1	15	1400	17	91	38.7
Dp	6	102	9600	13	136	28.9
Ds	24	374	35000	17	150	45.3
Dt	30	531	49000	17	150	64.3
Pr	5	47	4200	15	76	38.9
Pt	3	46	4300	17	65	31.0
Px	8	165	15700	17	150	49.8
Tf	7	57	5000	17	150	87.5

In areas of the Dillhut-Solvay complex, 0 to 3 percent slopes (map unit Ds), the average depth to the paleosol was 45.3 inches. One-half the observations had depths to paleosol between 22 and 62 inches. In the transected area of this map unit, depths to the paleosol were shallow at 21 percent, moderately deep at 37 percent, deep at 16 percent, and very deep at 26 percent of the observation points (see Figure 3, upper).

In areas of the Dillwyn-Tivoli fine sands, 0 to 15 percent slopes (map unit Dt), the average depth to the paleosol was 64.3 inches. One-half the observations had depths to paleosol between 29 and 86 inches. In the transected area of this map unit, depths to paleosol were shallow at 14 percent, moderately deep at 23 percent, deep at 17 percent, and very deep at 46 percent of the observation points (see Figure 3, lower).

In areas of Pratt loamy fine sand, 1 to 5 percent slopes (map unit Pr), the average depth to the paleosol was 38.9 inches. One-half the observations had depths to paleosol between 26 and 51 inches. In the transected area of this map unit, depths to paleosol were shallow at 17 percent, moderately deep at 43 percent, deep at 26 percent, and very deep at 14 percent of the observation points (see Figure 4, upper).

In areas of the Turon-Solvay-Carwile complex, 0 to 5 percent slopes (map unit Pt), the average depth to the paleosol was 31.0 inches. One-half the observations had depths to paleosol between 17 and 38

inches. In the transected area of this map unit, depths to paleosol were shallow at 33 percent, moderately deep at 43 percent, deep at 20 percent, and very deep at 4 percent of the observation points (see Figure 4, lower).

In areas of Langdon-Tivoli fine sands, 0 to 15 percent slopes (map unit Px), the average depth to the paleosol was 49.8 inches. One-half the observations had depths to paleosol between 17 and 64 inches. In the transected area of this map unit, depths to paleosol were shallow at 27 percent, moderately deep at 22 percent, deep at 22 percent, and very deep at 29 percent of the observation points (see Figure 5, upper).

In areas of the Tivoli fine sand, 10 to 30 percent slopes (map unit Tf), the average depth to the paleosol was 87.5 inches. One-half the observations had depths to paleosol between 55 and 123 inches. In the transected area of this map unit, depths to paleosol were shallow at 2 percent, moderately deep at 12 percent, deep at 14 percent, and very deep at 72 percent of the observation points (see Figure 5, lower).

### Systematic Sampling with EM Techniques

#### Introduction:

One of the most common uses for electromagnetic induction (EM) techniques has been to assess soil salinity (2, 3, 4, 5, 16, 17, 19) and sodicity (1). These techniques, however, have also been used to map bedrock surfaces (21), glacial deposits (20, 15), and permafrost (18, 9); estimate thicknesses of clays (14) or sand and gravel deposits (13); measure soil water content (11); and evaluate edaphic properties important to forest site productivity (12). These studies have documented some of the advantages of using EM interpretations and its applications over broad areas and soil types.

Electromagnetic induction techniques are a non-invasive geophysical method in which electromagnetic energy is used to measure the apparent conductivity of earthen materials. Apparent conductivity is the weighted average conductivity measurement for a column of earthen materials to a specified investigation depth (6). The averages are weighted according to the depth response function of the meter (19). Values of apparent conductivity are expressed in milliSiemens per meter (mS/m).

Apparent conductivity values are seldom diagnostic in themselves. However, lateral and vertical variations in these measurements can be used to infer changes in soils and soil properties. Electromagnetic responses are dependent on soil properties. Variations in electromagnetic responses are produced by changes in soil moisture, salt content, texture, and mineralogy (2). In nonsaline areas, soil texture, mineralogy, and moisture are the principal factors determining apparent conductivity (11). Each of these factors will affect the apparent conductivity of soils.

Soils and soil map units can be differentiated by characteristic ranges of apparent conductivity (7, 8). As EM measurements integrate several soil properties, responses can be correlated within geographic areas to soil types. Generally, in geophysical research, the term "soil type" has been used to refer to the particle-size class of unconsolidated sediments. Zalasiewicz and Mathers (20) used

EM techniques to map glacial sediments and distinguish areas of bouldery clays from arenaceous deposits. Sartorelli and French (18) defined grouping of the Unified Soil Classification System based on characteristic apparent conductivity values. Kingston (10) used EM techniques to distinguish and map soil and geologic deposits and strata. In these studies different "soil types" and lithologies induced distinct and characteristic EM responses.

Electromagnetic induction is an imperfect geophysical tool and is not equally suitable for use in all soil investigations. The success of an EM survey depends on the nature and variability of soil properties. This method has been most effective in areas where subsurface soil properties are fairly homogeneous, the effects of one factor dominant over the others, and variations in the EM response can be related to changes in a single factor (e.g. volumetric moisture, soluble salt content, clay content, soil depth, or mineralogy).

Each soil will have a characteristic EM response. For a particular soil, the EM response will constitute a range of values which will be influenced by temporal variations in soil moisture and temperature. Furthermore, cultural and natural features can be expected to influence these ranges. Similar soils will have a similar range in EM responses. Within a given geographic area, the conductivities of some soils will overlap. Soil properties and types can be inferred with EM techniques provided one is cognizant of changes in parent materials, drainage, topography, and vegetation.

Area of Langdon-Tivoli fine sands, 0 to 15 percent slopes.

Study Area:

The study site consisted of 3.2 acres of rangeland in southwestern Reno County, Kansas. Relief was about 10.6 feet. The area consists of several low mounds and ridges of windblown sands. Most areas are nearly level to rolling. The study site was bounded on the south by a fence line and road.

The topography of study site has been simulated in Figures 6 and 7. Figure 6 is a two-dimensional contour plot and Figure 7 is a three-dimensional surface net diagram of the ground surface. In Figure 6, the contour interval is 1.0 foot. The line of spot symbols appearing in Figure 6 represents the location of a radar traverse. In Figure 7, the vertical exaggeration is 10.

The study sites contains an intricate pattern of soils. Soils identified within the study site include Langdon (mixed, thermic Argic Ustipsamment), Tivoli (mixed, thermic Typic Ustipsamment), and Turon (sandy, mixed, thermic Psammentic Haplustalf).

Within the study site, at the time of this investigation, it was presumed that variations in the electromagnetic response would be directly related to changes in the depth to paleosol or lamellae. It was assumed that EM responses would be lower in areas of Tivoli and Langdon soils than in areas of Turon soil. In areas of soils with similar particle-size distributions, variations in the EM response would be highly influenced by differences in the volumetric moisture contents, and the presence, thickness, and arrangement of lamellae and soil horizons.

#### Field Methods:

A 400 by 350 foot rectangular grid was established across the study site. Survey flags were inserted in the ground at 50 foot intervals. At each of the 72 grid intersections, measurements were obtained with an EM38 meter in both the horizontal and vertical dipole orientations.

A transit was used to establish grid lines and determine the surface elevation of each grid intersection. Elevations were not tied to a benchmark; the lowest recorded surface point was chosen as the 0.0 foot datum.

#### Results:

Electromagnetic induction methods focuses on the rate and magnitude of change in EM response from place to place. Isarithmic maps prepared from EM data provide a graphic description of variations in soils and/or soil properties within a survey site.

Figures 8 and 9 represent two-dimensional isarithmic maps prepared from data collected with the EM38 meter in the horizontal and the vertical dipole orientations, respectively. The horizontal measurements integrates electromagnetic conductivity within the 0 to 0.75 m; the vertical measurements integrates values within the 0 to 1.5 m. At all observation points, values of apparent conductivity increased with increasing depth (horizontal dipole measurements < vertical dipole orientation). This relationship was believed to reflect increasing clay and volumetric moisture content with increasing soil depth.

Values of apparent conductivity were highest in the swales. These areas were inferred to have soils with higher clay contents (shallower depths to paleosol) and/or greater volumetric moisture contents. Values of apparent conductivity were lowest on the summits of dunes where the clay and volumetric moisture contents were lower. Intermediate values occurred on sideslopes.

A preliminary assessment of EM values was made in the field for the purpose of locating sampling sites. Based on this assessment, 5 observation points were selected and described. Table 3 list the soils and their averaged apparent conductivity values. Some soils can be differentiated by a characteristic range of apparent conductivity values. In Table 3, each soil appears to have a fairly unique EM response.

**Table 3**

**EM38 TRANSECT DATA  
AVERAGE APPARENT CONDUCTIVITY  
(in mS/m)**

<u>Soil</u>	<u># Observations</u>	<u>EMH</u>	<u>EMV</u>
Langdon	1	4.0	7.0
Tivoli	2	2.0	5.0
Turon	2	7.5	15.0

Using EM responses, the distribution of soils described at the five observation points was extended to other portions of the survey site.

Measurements taken with the EM38 meter in the vertical dipole orientation were used to prepare a soil map (Figure 10). Observation points with low ( $< 6$  mS/m), intermediate (6 to 11 mS/m), and high ( $> 11$  mS/m) EM responses were identified as Tivoli, Langdon, and Turon soils, respectively.

A radar traverse was completed along the grid line identified by the spot symbols in figures 6, 8, and 9. The enclosed radar profile is from this grid line. It has been processed through RADAN software. The profile has been normalized, terrain corrected, and the color tables and values have been transformed.

The numbers appearing at the top of the radar profile represent distances (in feet) and grid intersections. A paleosol (P) is evident beneath the dune in the left-hand portion of the radar profile. This layer varies in depth from about 3.0 feet in a swale near the 400 foot mark to about 7.5 feet on the summit of the dune near the 300 foot mark. Depths to this interface are moderately deep between the 150 and 50 foot marks. The strong and reverberated signals beneath "M" in the left-hand portion of the radar profile were produced by a buried metallic object.

#### Area of Tabler-Punkin complex.

##### Survey Area:

The study site consisted of about 3.6 acres of cropland in northeastern Reno County, Kansas. Relief was about 1.4 feet. The study site was bounded on the north by a road. Tabler is a member of the fine, montmorillonitic, thermic Udertic Argiustoll family. The proposed Punkin soil is a member of the fine, mixed, thermic Aquic Natrustoll family.

##### Field Methods:

A 350 by 450 foot rectangular grid was established across the study site. Survey flags were inserted in the ground at 50 foot intervals. At each of the 80 grid intersections, measurements were obtained with an EM38 meter in both the horizontal and vertical dipole orientations.

A transit was used to establish grid lines and determine surface elevations at each grid intersect. Elevations were not tied to a benchmark; the lowest recorded surface point was chosen as the 0.0 foot datum.

##### Results:

The topography of study site has been simulated in Figures 11 and 12. Figure 11 is a two-dimensional contour plot and Figure 12 is a three-dimensional surface net diagram of the ground surface. In Figure 11, the contour interval is 0.2 foot. In Figure 12, the vertical exaggeration is 20. This scale and vertical exaggeration were needed to express an old field-boundary line in the south-central portion of the study area.

Figures 13 and 14 represent two dimensional plots of apparent conductivity values collected in the horizontal and vertical dipole orientations, respectively. In each figure, the interval is 10 mS/m.

For all but five observation points, EM response increased with increasing soil depth (response in the vertical greater than in the

conductive materials (i.e. greater clay, soluble salt, and/or water contents) with increasing soil depth. The reverse relationship at several observation points was attributed to high concentrations of soluble salts near the soil surface.

In both plots, EM responses appear to be highly variable across the site. Responses decreased toward the right-hand margin of each plot indicating a probable reduction in the amount of soluble salts and a change in soil type. In addition, in both plots, a noticeable linear pattern extends across the site from the northwest to the southeast corners. This feature may reflect a buried fluvial feature which exerts control over the soils.

Electromagnetic responses were relatively high across this site. Based on 80 observations, the average apparent conductivity was 103 mS/m, with a range of 27 to 180 mS/m in the horizontal mode. One-half of the observations collected in the horizontal dipole orientation had apparent conductivity values between 72 and 129 mS/m. The average apparent conductivity was 127 mS/m, with a range of 42 to 211 mS/m in the vertical mode. One-half of the observations collected in the vertical dipole orientation had apparent conductivity values between 96 and 158 mS/m.

#### **Conclusions:**

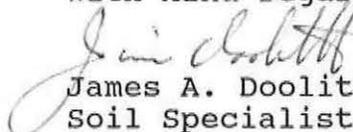
1. Ground-penetrating radar techniques are not appropriate for use in many soils of Reno County. Depths of penetration were restricted by thick bands of lamellae or layers of coarse-loamy or finer textured soil materials. In areas of coarse-textured soil materials, GPR was a suitable tool and provided adequate depths of penetration and good resolution of subsurface interfaces. However, in many areas of coarse-textured materials, while the depth of observation and resolution of GPR was good, the large number and segmented nature of subsurface layers plagued interpretations. To properly identify each layer, a prohibitive number of auger observations would have been required. One layer, the "paleosol", was distinct and recognizable across most radar profiles. Ground-penetrating radar techniques were used to characterize the depth to "paleosols" across extensive areas and several map units.

2. Continuous or discrete sampling with EM meters along transect lines can be used to detect variations in soils and soil properties. This tool can be used to provide objective description of soils, evaluate "modal" soil conditions, locate some map unit boundaries, and identify sites for more detailed observations. In Reno County, the EM38 meter appears to be a useful tool for assessing sodium-affected soils and map units. All participants received training and exposure to the uses of the EM38 meter. The NSSC has loaned an EM38 meter (serial number 8906008) to the survey party for the period of 1 August to 1 November 1994.

3. Isarithmic maps prepared from EM data can provide a graphic description of variations in soils and/or soil properties within selected map units. I encourage the soil survey party to avail themselves to these techniques.

It was my pleasure to work with and assist the soil scientists of Reno County.

With kind regards

  
James A. Doolittle  
Soil Specialist

cc:

James Culver, National Leader, SSQAS, NSSC, SCS, Lincoln, NE  
Paul Finnell, Assistant State Soil Scientist, SCS, Salina, KS  
Steve Holzhey, National Leader, SSQAS, NSSC, SCS, Lincoln, NE  
Robert Murphy, Soil Party Leader, SCS, Hutchinson Area Office,  
West 28th Street, Suite B, Hutchinson, KS, 67502

## REFERENCES

1. Ammons, J. T., M. E. Timpson, and D. L. Newton. 1989. Application of aboveground electromagnetic conductivity meter to separate Natraqualfs and Ochraqualfs in Gibson County, Tennessee. *Soil Survey Horizons* 30(3):66-70.
2. Cook, P. G. and G. R. Walker. 1992. Depth profiles of electrical conductivity from linear combinations of electromagnetic induction measurements. *Soil Sci. Soc. Am. J.* 56:1015-1022.
3. Corwin, D. L., and J. D. Rhoades. 1982. An improved technique for determining soil electrical conductivity-depth relations from above-ground electromagnetic measurements. *Soil Sci. Soc. Am. J.* 46:517-520.
4. Corwin, D. L., and J. D. Rhoades. 1984. Measurements of inverted electrical conductivity profiles using electromagnetic induction. *Soil Sci. Soc. Am. J.* 48:288-291.
5. Corwin, D. L., and J. D. Rhoades. 1990. Establishing soil electrical conductivity - depth relations from electromagnetic induction measurements. *Communications in Soil Sci. Plant Anal.* 21(11&12):861-901.
6. Greenhouse, J. P., and D. D. Slaine. 1983. The use of reconnaissance electromagnetic methods to map contaminant migration. *Ground Water Monitoring Review* 3(2):47-59.
7. Hoekstra, P., R. Lahti, J. Hild, R. Bates, and D. Phillips. 1992. Case histories of shallow time domain electromagnetics in environmental site assessments. *Ground Water Monitoring Review.* 12(4):110-117.
8. Jaynes, D. B., T. S. Colvin, J. Ambuel. 1993. Soil Type and crop yield determination from ground conductivity surveys. 1993 International Meeting of American Society of Agricultural Engineers. Paper No. 933552. ASAE, St. Joseph, MI. pp. 6.
9. Kawasaki, K. and T. E. Osterkamp. 1988. Mapping shallow permafrost by electromagnetic induction - practical considerations. *Cold Region Science and Technology* 15: 279-288.
10. Kingston, G. 1985. Electromagnetic inductive instruments for use in surveys of soil salinity. *Proc. Australian Soc. Sugar Cane Tech.* Brisbane, Queensland. p. 74-84.
11. Kachanoski, R. G., E. G. Gregorich, and I. J. Van Wesenbeeck. 1988. Estimating spatial variations of soil water content using noncontacting electromagnetic inductive methods. *Can. J. Soil Sci.* 68:715-722.
12. McBride, R. A., A. M. Gordon, and S. C. Shrive. 1990. Estimating forest soil quality from terrain measurements of apparent electrical conductivity. *Soil Sci. Soc. Am. J.*, 54:290-293.
13. McNeill, J. D. 1988. Advances in electromagnetic methods for groundwater studies. Geonics Ltd. Mississauga, Canada.

14. Palacky, G. J. 1987. Clay mapping using electromagnetic methods. *First Break* 5(8):295-306.
15. Palacky, G. J. and L. E. Stephens. 1990. Mapping of Quaternary sediments in northeastern Ontario using ground electromagnetic methods. *Geophysics* 55(12):1596-1604.
16. Rhoades, J. D. and D. L. Corwin. 1981. Determining soil electrical conductivity-depth relations using an inductive electromagnetic soil conductivity meter. *Soil Sci. Soc. Am. J.* 45:255-260.
17. Rhoades, J. D., N. A. Manteghi, P. J. Shouse, and W. J. Alves. 1989. Soil Electrical conductivity and soil salinity: new formulation and calibrations. *Soil Sci. Soc. Am. J.* 53:433-439.
18. Sartorelli, A. N. and R. B. French. 1982. Electro-magnetic induction methods for mapping permafrost along northern pipeline corridors. pp. 283-295. In: French, H. M. (ed.) *Proceedings of the Fourth Canadian Permafrost Conference*. Calgary, Alberta 2-6 March 1991. National Research Council of Canada.
19. Wollenhaupt, N. C., J. L. Richardson, J. E. Foss, and E. C. Doll. 1986. A rapid method for estimating weighted soil salinity from apparent soil electrical conductivity measured with an aboveground electromagnetic induction meter. *Can J. Soil Sci.* 66:315-321.
20. Zalasiewicz, J. A. and S. J. Mathers. 1985. New approach to mapping tills. pp. 55-59. IN: M. C. Forde (ed.) *Proceedings of the International Conference on Construction in Glacial Tills and Boulder Clays*. 12-14 March 1985. Edinburgh. (n. p. Engineering Technics Press).
21. Zalasiewicz, J. A., S. J. Mathers, and J. D. Cornwell. 1985. The application of ground conductivity measurements to geological mapping. *Q. J. English Geol. London* 18:139-148.

## ADDENDUM

Transect data on the depth to paleosol collected with GPR in Reno County, Kansas:

TRANSECT 1A		TRANSECT 1		TRANSECT 2	
MU: DT		MU: DP		MU: DT	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	79"	0'	54"	0'	36"
100	79	100	58	100	34
200	74	200	57	200	32
300	68	300	44	300	33
400	56	400	62	400	34
500	57	500	57	500	74
600	28	600	40	600	36
700	29	700	37	700	34
800	28	800	46	800	40
900	29	900	36	900	83
		1000	37	1000	19
		1100	31	1100	96
		1200	32		
		1300	35		
		1400	17		
		1500	27		
		1600	29		
		1700	34		

TRANSECT 3		TRANSECT 4		TRANSECT 5	
MU: DT		MU: TF		MU: DT	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	45"	0'	26"	0'	80"
100	54	100	92	100	27
200	45	200	106	200	34
300	27	300	70	300	17
400	45			400	17
500	17			500	28
600	17			600	62
700	40				
800	45				
900	31				
1000	17				
1100	24				
1200	17				

TRANSECT 6		TRANSECT 7		TRANSECT 8	
MU: DT		MU: DP		MU: PR	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	23"	0'	34"	0'	26"
100	17	100	33	100	15
200	17	200	17		
300	17	300	17		
400	17	400	17		
500	17	500	17		
600	35	600	25		
		700	24		
		800	25		
		900	26		
		1000	17		
		1100	17		
		1200	17		
		1300	17		
		1400	17		
		1500	17		
		1600	17		
		1700	17		
		1800	17		
		1900	17		
		2000	17		
		2100	17		
		2200	41		
		2300	25		
		2400	13		
TRANSECT 9		TRANSECT 10		TRANSECT 11	
MU: DP		MU: DT		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	17"	0'	25"	0'	17"
100	17	100	54	100	17
200	17	200	54	200	17
300	17	300	32	300	17
400	17	400	19	400	17
500	17	500	31	500	17
600	17	600	28	600	17
700	17	700	17	700	17
800	56	800	17	800	17
900	73	900	17	900	17
1000	49	1000	17	1000	19
1100	17	1100	17	1100	37
1200	17	1200	17	1200	68
1300	17	1300	20	1300	118
1400	17	1400	49	1400	35
1500	17	1500	25	1500	34
1600	17	1600	19	1600	18
1700	17	1700	29	1700	69
1800	17	1800	32	1800	32
1900	17	1900	17	1900	31
2000	41	2000	17	2000	45
2100	29	2100	35	2100	17
2200	34	2200	17	2200	17
2300	25	2300	27	2300	17
2400	17	2400	27	2400	17
2500	17	2500	34	2500	36
2600	17			2600	35

TRANSECT 12		TRANSECT 13		TRANSECT 14	
MU: PR		MU: DT		MU: DT	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	17"	0'	28"	0'	17"
100	17	100	37	100	25
200	37	200	28	200	35
300	59	300	37	300	23
400	75	400	37	400	34
500	65	500	33	500	52
600	76	600	17	600	71
700	70	700	31	700	53
800	42	800	17	800	29
900	45	900	17	900	17
1000	49	1000	37	1000	17
1100	28	1100	43	1100	17
1200	62	1200	17	1200	51
1300	51	1300	17	1300	62
1400	31	1400	80	1400	51
1500	32	1500	80	1500	52
1600	17	1600	80	1600	45
1700	37	1700	80	1700	34
1800	40	1800	80	1800	38
1900	17	1900	84	1900	39
2000	27	2000	52	2000	45
2100	23	2100	41	2100	67
2200	17	2200	41		
2300	34	2300	28		
2400	32	2400	34		
2500	32	2500	52		
		2600	68		

TRANSECT 15		TRANSECT 16		TRANSECT 17	
MU: DT		MU: PR		MU: DT	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	37"	0'	32"	0'	71"
100	28	100	25	100	91
200	26	200	29	200	90
300	34	300	18	300	42
400	17	400	23	400	43
500	36	500	33	500	77
600	56	600	51	600	34
700	77	700	61	700	50
800	95	800	58	800	28
900	69			900	17
1000	60			1000	100
1100	56			1100	121
1200	51			1200	72
1300	46			1300	67
1400	41			1400	127
1500	49			1500	104
1600	69			1600	51
1700	40			1700	91
1800	34			1800	82
1900	27			1900	40
2000	24				
2100	31				
2200	17				
2300	17				
2400	17				
2500	26				
2600	17				
2700	72				
2800	45				

TRANSECT 18		TRANSECT 19		TRANSECT 20	
MU: DT		MU: TF		MU: DT	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	54"	0'	34"	0'	33"
100	85	100	79	100	18
200	73	200	48	200	17
300	68			300	75
400	50			400	36
500	68			500	29
600	60			600	17
700	57			700	17
800	72			800	17
900	79			900	17
1000	34			1000	17
1100	20				
1200	56				
1300	99				
1400	17				
1500	43				
1600	91				
1700	29				
1800	17				
1900	63				
2000	134				
2100	85				
2200	17				
2300	105				
2400	100				
2500	38				
2600	150				
2700	150				
2800	150				
2900	150				
3000	150				
3100	77				

TRANSECT 21		TRANSECT 22		TRANSECT 23	
MU: DP		MU: DT		MU: DP	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	23"	0'	65"	0'	57"
100	29	100	120	100	76
200	21	200	150	200	69
300	29	300	150	300	136
400	17	400	150	400	85
500	17	500	150		
600	17	600	88		
700	18	700	116		
800	26				
900	34				
1000	32				
1100	31				
1200	63				

TRANSECT 24		TRANSECT 25		TRANSECT 26	
MU: TF		MU: DT		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	68"	0'	28"	0'	84"
100	87	100	37	100	85
200	150	200	93	200	83
300	150	300	94	300	48
400	104			400	70
				500	54
				600	36
				700	36
				800	37
				900	32
				1000	63

TRANSECT 27		TRANSECT 28		TRANSECT 29	
MU: DT		MU: TF		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	46"	0'	120"	0'	150"
100	93	100	150	100	150
200	127	200	138	200	150
300	150	300	130		
400	150				
500	150				
600	150				
700	136				
800	150				
900	88				

## TRANSECT 27

MU: DT

DISTANCE	DEPTH
0'	46"
100	93
200	127
300	150
400	150
500	150
600	150
700	136
800	150
900	88

## TRANSECT 28

MU: TF

DISTANCE	DEPTH
0'	120"
100	150
200	138
300	130

## TRANSECT 29

MU: DS

DISTANCE	DEPTH
0'	150"
100	150
200	150

## TRANSECT 30

MU: TF

DISTANCE	DEPTH
0'	84"
100	119

## TRANSECT 31

MU: DS

DISTANCE	DEPTH
0'	135"
100	123

## TRANSECT 32

MU: DP

DISTANCE	DEPTH
0'	17"
100	29
200	18
300	29
400	25
500	17
600	17
700	17
800	17
900	17
1000	17
1100	17
1200	17
1300	17

## TRANSECT 33

MU: PR

DISTANCE	DEPTH
0'	48"
100	52
200	61
300	41

## TRANSECT 34

MU: DS

DISTANCE	DEPTH
0'	33"
100	20
200	18
300	21
400	32

## TRANSECT 35

MU: DS

DISTANCE	DEPTH
0'	26"
100	17
200	17
300	17
400	17
500	17
600	17
700	17
800	21
900	25
1000	28
1100	32

**TRANSECT 36**

MU: DT

<u>DISTANCE</u>	<u>DEPTH</u>
0'	44"
100	57
200	24
300	29
400	29
500	41
600	49
700	91
800	37
900	79
1000	49
1100	117
1200	43

**TRANSECT 37**

MU: DS

<u>DISTANCE</u>	<u>DEPTH</u>
0'	25"
100	36
200	23
300	18
400	17
500	17
600	44
700	17
800	41
900	57
1000	28
1100	40
1200	43
1300	33
1400	57
1500	85
1600	87

**TRANSECT 38**

MU: DS

<u>DISTANCE</u>	<u>DEPTH</u>
0'	23"
100	20
200	17
300	23
400	28
500	43
600	17
700	17
800	17
900	23
1000	54
1100	54
1200	37

**TRANSECT 39**

MU: PT

<u>DISTANCE</u>	<u>DEPTH</u>
0'	42"
100	50
200	42
300	39
400	17
500	17
600	17
700	37
800	17
900	17
1000	17
1100	22
1200	29
1300	28
1400	34
1500	36
1600	27
1700	43
1800	36
1900	33
2000	48

**TRANSECT 40**

MU: DS

<u>DISTANCE</u>	<u>DEPTH</u>
0'	24"
100	23
200	25
300	23
400	25
500	23
600	22
700	24
800	29
900	28
1000	32
1100	25

**TRANSECT 41**

MU: DS

<u>DISTANCE</u>	<u>DEPTH</u>
0'	53"
100	58
200	62
300	31
400	17
500	31
600	31
700	57
800	44
900	56
1000	76
1100	76
1200	44

TRANSECT 42		TRANSECT 43		TRANSECT 44	
MU: DT		MU: DS		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	34"	0'	45"	0'	78"
100	29	100	48	100	51
200	29	200	42	200	53
300	62	300	59	300	37
400	85	400	61	400	74
500	75	500	77	500	57
600	28	600	91	600	75
700	28	700	75	700	61
800	28	800	71	800	48
900	28	900	84	900	52
1000	34	1000	52	1000	62
1100	36	1100	83	1100	53
		1200	91	1200	48
		1300	85	1300	50
		1400	74	1400	28
		1500	73	1500	17
		1600	94		
		1700	110		

TRANSECT 45		TRANSECT 46		TRANSECT 47	
MU: DT		MU: DS		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	150"	0'	53"	0'	18"
100	150	100	32	100	23
200	125	200	34	200	21
300	77	300	28	300	21
400	84	400	35	400	24
500	126	500	32	500	23
600	68	600	28	600	22
700	79	700	49	700	26
800	62	800	92	800	20
900	125	900	49	900	42
1000	119	1000	17	1000	65
1100	84	1100	17	1100	36
1200	74	1200	17	1200	35
1300	113	1300	47	1300	21
1400	91	1400	57	1400	23
1500	125	1500	70	1500	32
1600	136			1600	26
1700	96			1700	24
1800	35			1800	21
1900	150			1900	22
2000	51			2000	24
2100	42			2100	23
2200	150			2200	22
2300	51			2300	22
2400	42				
2500	33				
2600	39				
2700	36				
2800	62				
2900	65				
3000	108				
3100	75				
3200	113				
3300	99				
3400	49				
3500	82				
3600	84				
3700	74				
3800	65				
3900	31				
4000	57				
4100	108				
4200	82				
4300	123				
4400	150				
4500	150				
4600	86				
4700	150				
4800	150				

TRANSECT 48		TRANSECT 49		TRANSECT 50	
MU: DT		MU: DT		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	67"	0'	23"	0'	25"
100	96	100	79	100	24
200	66	200	45	200	26
300	77	300	61	300	40
400	46	400	86	400	19
500	53	500	24	500	22
600	49	600	25	600	23
700	37	700	26	700	24
800	45	800	25	800	17
900	45			900	17
1000	69				
1100	96				
1200	86				
1300	113				
1400	100				
1500	87				
1600	126				
1700	62				
1800	60				
1900	150				
2000	83				
2100	70				
2200	49				
2300	74				
2400	87				
2500	59				
2600	29				
2700	37				
2800	137				
2900	57				
3000	133				
3100	85				
3200	97				
3300	150				
3400	150				
3500	150				
3600	150				

## TRANSECT 51

MU: DT

DISTANCE	DEPTH
0'	29"
100	23
200	27
300	17
400	17
500	17
600	61
700	150
800	48
900	17
1000	17
1100	17
1200	17
1300	17
1400	17
1500	17
1600	20

## TRANSECT 52

MU: TF

DISTANCE	DEPTH
0'	26"
100	65
200	108
300	134
400	86
500	82
600	59
700	25
800	71
900	147
1000	150
1100	150
1200	150
1300	150

## TRANSECT 53

MU: DS

DISTANCE	DEPTH
0'	54"
100	37
200	49
300	150
400	150
500	150
600	91
700	64
800	91
900	46
1000	26

## TRANSECT 54

MU: DT

DISTANCE	DEPTH
0'	40"
100	31
200	71
300	62
400	76
500	68
600	105
700	150
800	150
900	150
1000	130
1100	97
1200	113
1300	117

## TRANSECT 55

MU: DT

DISTANCE	DEPTH
0'	54"
100	79
200	52
300	43
400	68
500	150
600	122
700	102
800	43
900	35
1000	26
1100	37
1200	49
1300	41
1400	102
1500	76

## TRANSECT 56

MU: TF

DISTANCE	DEPTH
0'	74"
100	94
200	65
300	43
400	74
500	86
600	150
700	150
800	95

TRANSECT 57		TRANSECT 58		TRANSECT 59	
MU: DT		MU: DS		MU: DS	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	130"	0'	36"	0'	40"
100	128	100	40	100	17
200	103	200	25	200	17
300	82	300	32	300	33
400	76	400	27	400	24
500	85	500	29	500	17
600	41	600	29	600	24
700	45	700	34	700	38
800	68	800	79	800	17
900	77	900	24	900	53
1000	45	1000	26	1000	27
1100	84	1100	17	1100	25
		1200	17	1200	63
		1300	17	1300	114
		1400	17	1400	78
		1500	17	1500	76
		1600	17	1600	128
		1700	51	1700	150
		1800	27		
		1900	26		
		2000	23		
		2100	22		
		2200	66		
		2300	31		
		2400	17		
		2500	37		
		2600	17		
		2700	17		

TRANSECT 60		TRANSECT 61		TRANSECT 62	
MU: DS		MU: DS		MU: DT	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	68"	0'	44"	0'	85"
100	73	100	22	100	65
200	67	200	31	200	68
300	80	300	23	300	56
400	74	400	17	400	108
500	63	500	17	500	150
600	57	600	74	600	150
700	77	700	42	700	150
800	88	800	21	800	150
900	69	900	36	900	150
1000	67	1000	38	1000	150
1100	94	1100	83	1100	150
1200	88	1200	78	1200	150
1300	97	1300	80	1300	150
1400	93	1400	41	1400	104
1500	78	1500	33	1500	59
1600	71	1600	20	1600	17
1700	60	1700	28	1700	31
1800	67	1800	62	1800	17
1900	96	1900	40	1900	42
2000	150	2000	17	2000	114
2100	150	2100	26	2100	150
2200	77	2200	19	2200	150
2300	63	2300	32	2300	129
2400	31	2400	41	2400	118
2500	25	2500	28	2500	82
2600	26	2600	17	2600	78
2700	28	2700	17	2700	26
2800	17	2800	50	2800	17
2900	17	2900	17	2900	28
3000	17	3000	49	3000	17
3100	17	3100	48	3100	17
3200	17	3200	53	3200	40
3300	27	3300	85	3300	29
3400	32	3400	27	3400	51
		3500	68	3500	17
		3600	45	3600	26
		3700	62	3700	32
		3800	51	3800	17
		3900	65	3900	17
				4000	54
				4100	58
				4200	83
				4300	47
				4400	57
				4500	42
				4600	17
				4700	17
				4800	17
				4900	62
				5000	84
				5100	93
				5200	76
				5300	150
				5400	62

**TRANSECT 63**

MU: DT

<u>DISTANCE</u>	<u>DEPTH</u>
0'	121"
100	66
200	51
300	79
400	72
500	84
600	100
700	39
800	50
900	57
1000	96
1100	101
1200	103
1300	78
1400	88
1500	87
1600	72

**TRANSECT 64**

MU: DS

<u>DISTANCE</u>	<u>DEPTH</u>
0'	25"
100	17
200	27
300	29
400	29
500	17
600	17
700	17
800	32
900	17
1000	17
1100	17
1200	17

**TRANSECT 65**

MU: TF

<u>DISTANCE</u>	<u>DEPTH</u>
0'	68"
100	68
200	150
300	103
400	53
500	50
600	127
700	70
800	17
900	25
1000	113
1100	50
1200	80
1300	68
1400	110
1500	150

**TRANSECT 66**

MU: DS

<u>DISTANCE</u>	<u>DEPTH</u>
0'	46"
100	28
200	56
300	51
400	50
500	50
600	34
700	34
800	109
900	104
1000	126
1100	97
1200	126
1300	88
1400	117

**TRANSECT 67**

MU: DT

<u>DISTANCE</u>	<u>DEPTH</u>
0'	150"
100	91
200	78
300	97
400	70
500	87

**TRANSECT 68**

MU: DT

<u>DISTANCE</u>	<u>DEPTH</u>
0'	40"
100	17
200	27
300	26
400	31
500	28
600	88
700	105
800	59
900	74
1000	70

## TRANSECT 69

MU: TF

DISTANCE	DEPTH
0'	28"
100	131
200	48
300	60
400	61
500	150

## TRANSECT 70

MU: DS

DISTANCE	DEPTH
0'	108"
100	70
200	27
300	24
400	34
500	58
600	71
700	72
800	34
900	73
1000	101

## TRANSECT 71

MU: DT

DISTANCE	DEPTH
0'	82"
100	103
200	86
300	75
400	61
500	74
600	104
700	46
800	71

## TRANSECT 72

MU: DS

DISTANCE	DEPTH
0'	51"
100	34
200	17
300	54

## TRANSECT 73

MU: DT

DISTANCE	DEPTH
0'	59"
100	150
200	102
300	150
400	150
500	150
600	73
700	29
800	66
900	150
1000	28
1100	51
1200	17
1300	36

## TRANSECT 74

MU: PX

DISTANCE	DEPTH
0'	17"
100	31
200	17
300	17
400	23
500	32
600	17
700	41
800	80
900	80
1000	80
1100	36
1200	60
1300	40
1400	42
1500	31
1600	44
1700	31
1800	108
1900	27
2000	77
2100	17
2200	17
2300	20
2400	17
2500	34
2600	34
2700	17
2800	17
2900	33
3000	38
3100	17
3200	17
3300	41
3400	26
3500	88
3600	17
3700	38
3800	40
3900	26

## TRANSECT 75

MU: PR

DISTANCE	DEPTH
0'	52"
100	35
200	17
300	52
400	37
500	31

## TRANSECT 76

MU: PX

DISTANCE	DEPTH
0'	100"
100	94
200	82
300	48

## TRANSECT 77

MU: PT

DISTANCE	DEPTH
0'	49"
100	51
200	23
300	17
400	28
500	27
600	20
700	57
800	28
900	34
1000	65
1100	23

## TRANSECT 78

MU: PX

DISTANCE	DEPTH
0'	42"
100	79
200	42
300	35
400	99
500	54
600	31

## TRANSECT 79

MU: PT

DISTANCE	DEPTH
0'	31"
100	44
200	17
300	17
400	26
500	17
600	34
700	35
800	17
900	17
1000	17
1100	17
1200	65

## TRANSECT 80

MU: AX

DISTANCE	DEPTH
0'	36"
100	22
200	17
300	45
400	47
500	36
600	91
700	32
800	40
900	36
1000	34
1100	34
1200	34
1300	34
1400	43

TRANSECT 81		TRANSECT 82		TRANSECT 83	
MU: PX		MU: TX		MU: PX	
DISTANCE	DEPTH	DISTANCE	DEPTH	DISTANCE	DEPTH
0'	40"	0'	150"	0'	17"
100	47	100	95	100	34
200	17	200	37	200	110
300	66	300	91	300	17
400	22	400	44	400	53
500	24	500	47	500	150
600	150	600	53	600	84
700	150	700	17	700	17
800	104	800	17	800	69
900	33	900	63	900	76
1000	61	1000	17	1000	17
1100	34	1100	17	1100	42
1200	28	1200	17	1200	150
1300	27	1300	17	1300	133
1400	27	1400	17	1400	26
1500	27	1500	43	1500	50
1600	44	1600	61	1600	71
1700	76	1700	58	1700	117
1800	54	1800	17	1800	150
1900	45	1900	76	1900	88
2000	17	2000	69		
2100	35	2100	26		
2200	17				
2300	36				
2400	35				

## TRANSECT 84

MU: PX

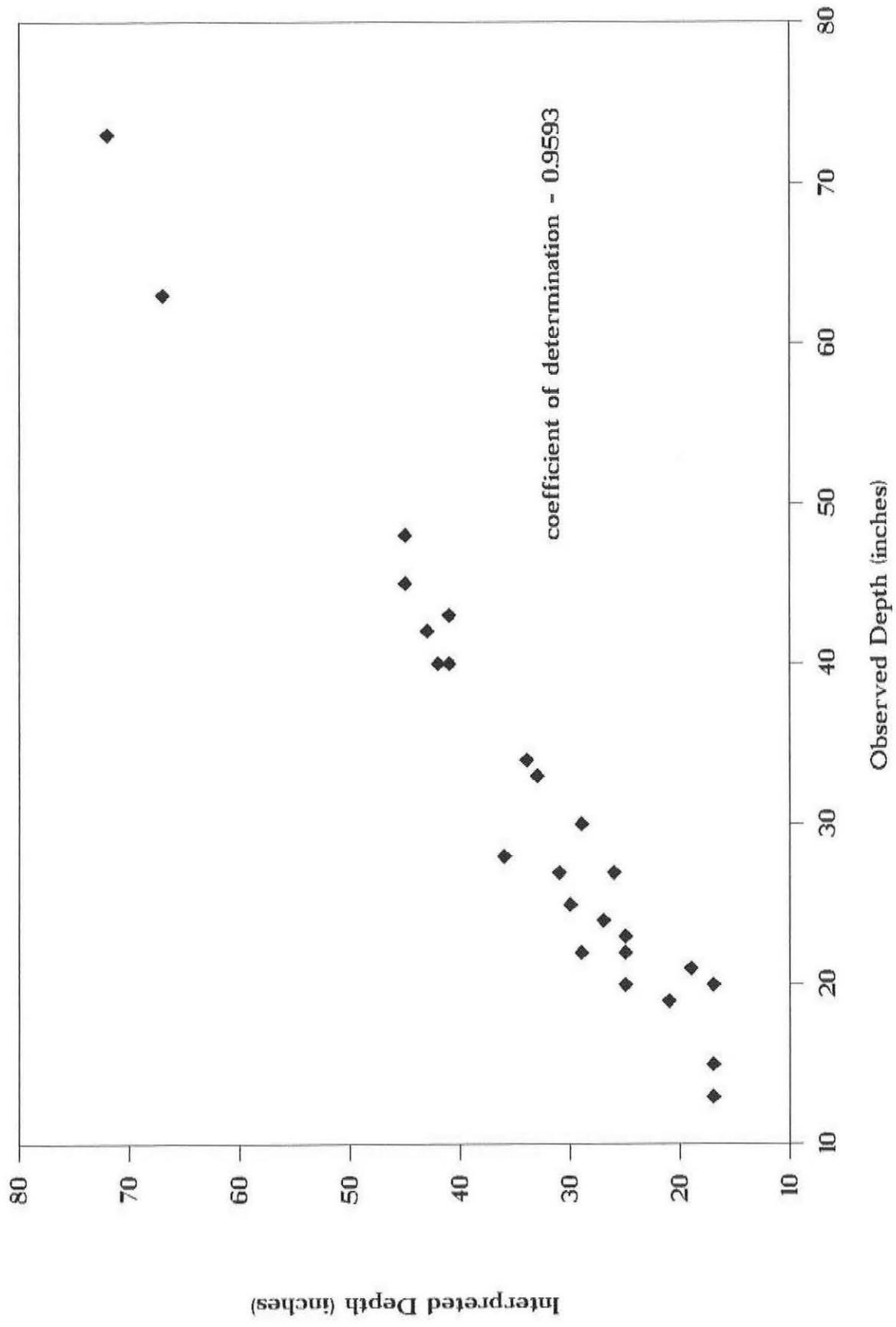
<u>DISTANCE</u>	<u>DEPTH</u>
0'	44"
100	17
200	26
300	41
400	83
500	99
600	33
700	52
800	17
900	57
1000	49
1100	17
1200	60
1300	60
1400	37
1500	17
1600	17
1700	17
1800	17
1900	45
2000	17
2100	59
2200	43
2300	57
2400	17
2500	136
2600	62
2700	17
2800	74
2900	50
3000	68
3100	118
3200	127
3300	57

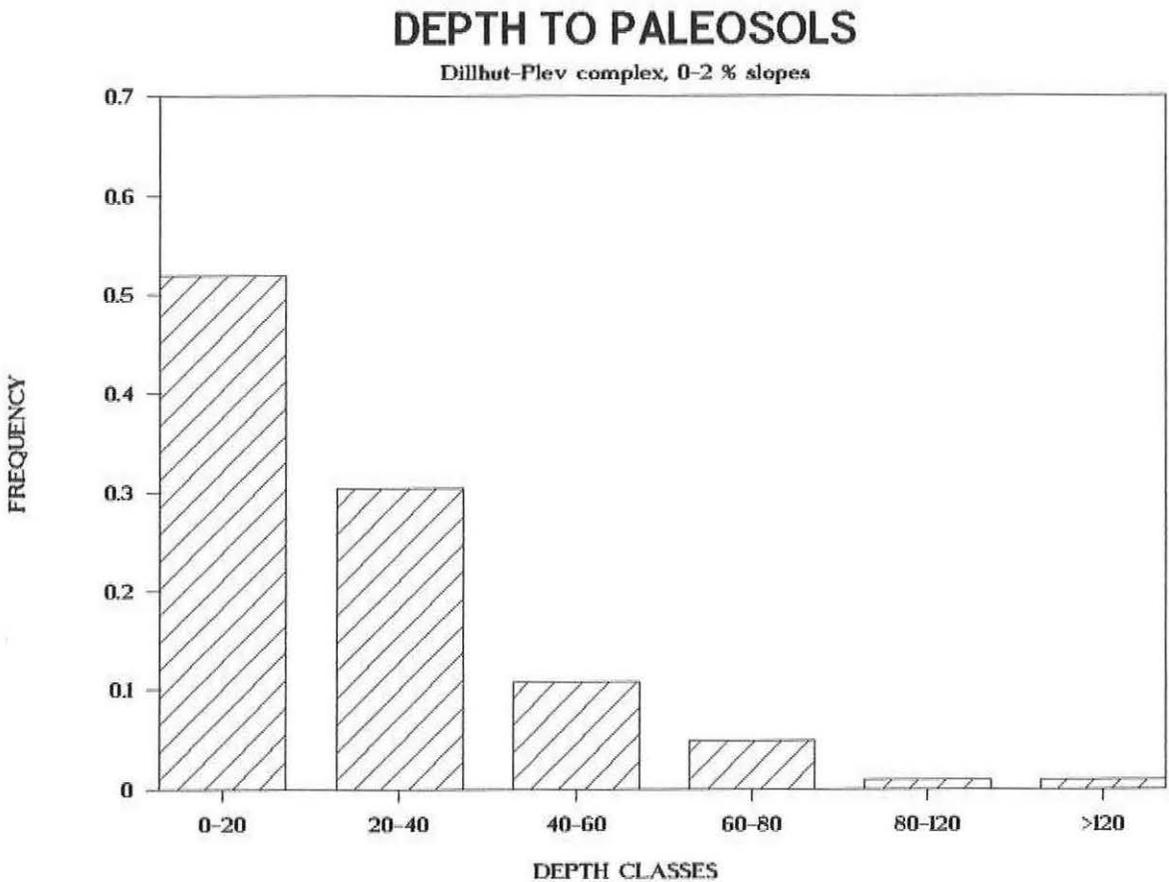
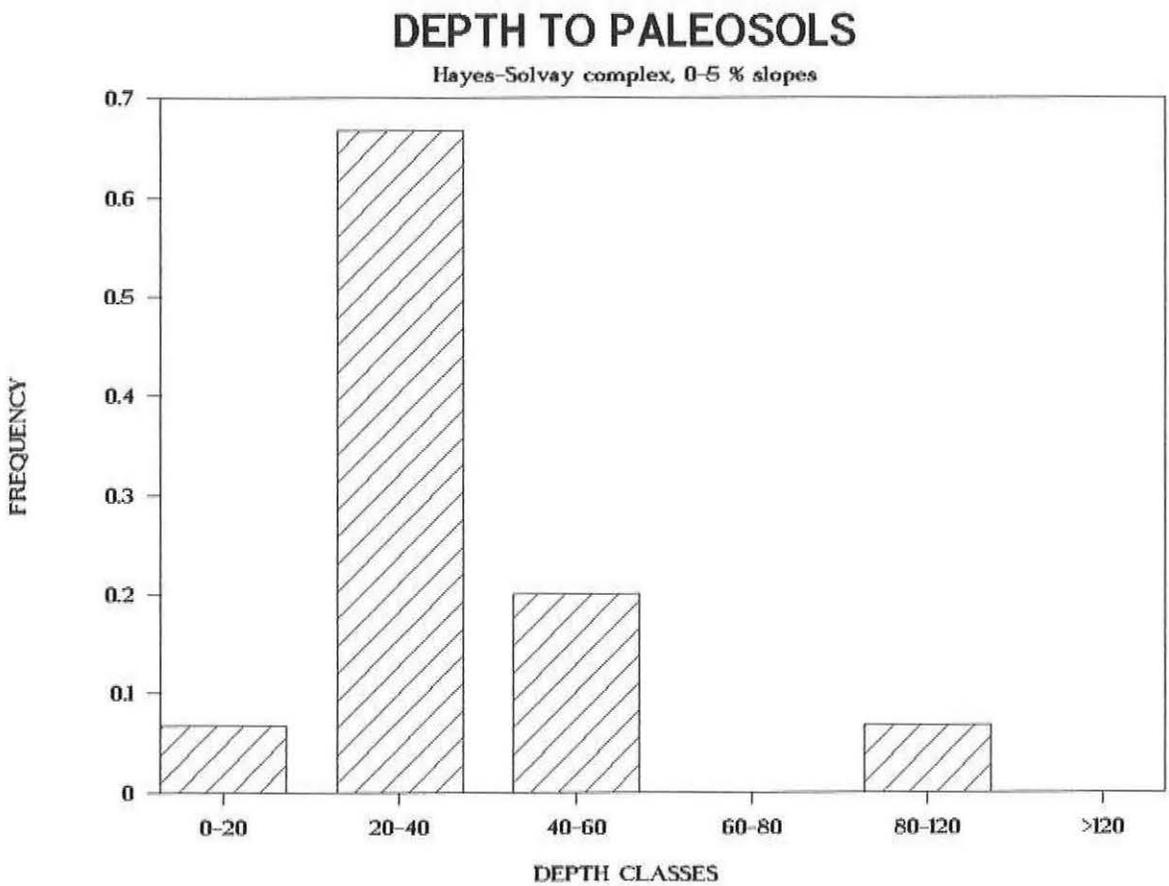
## TRANSECT 85

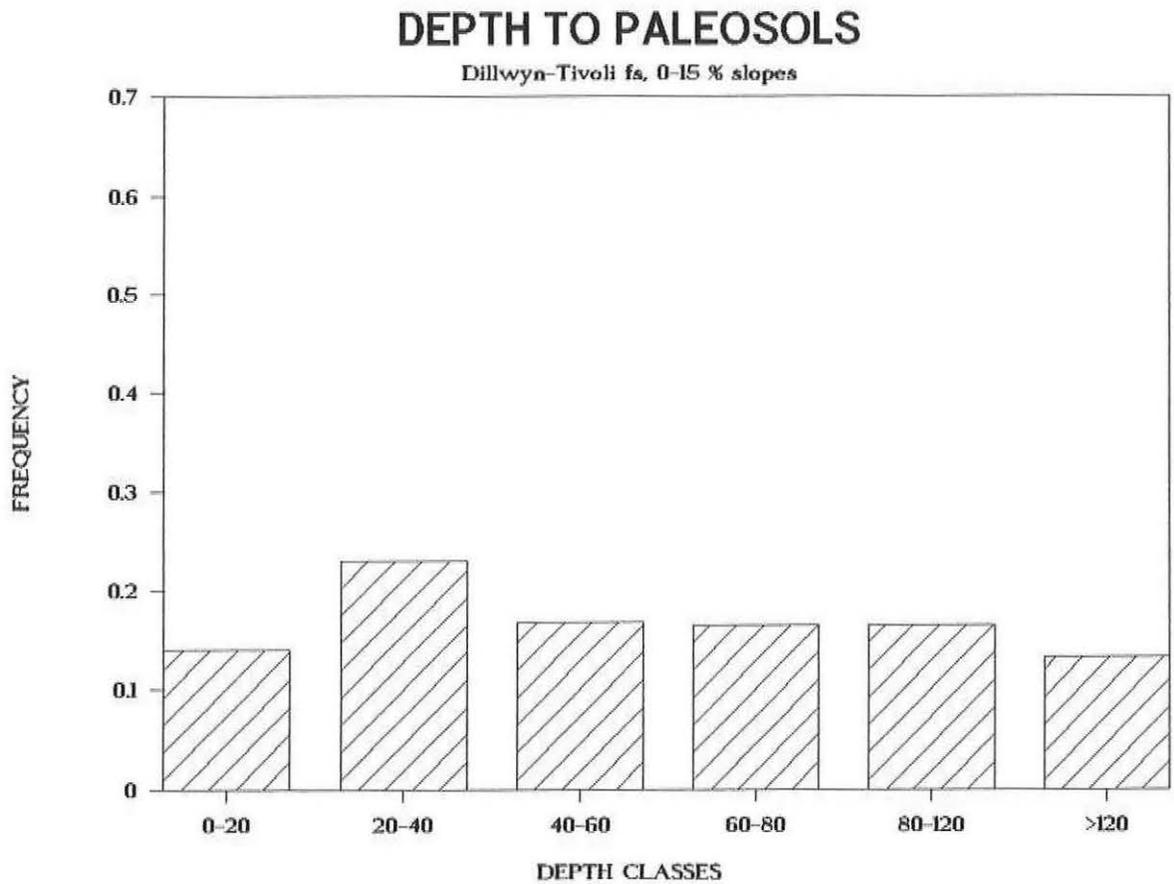
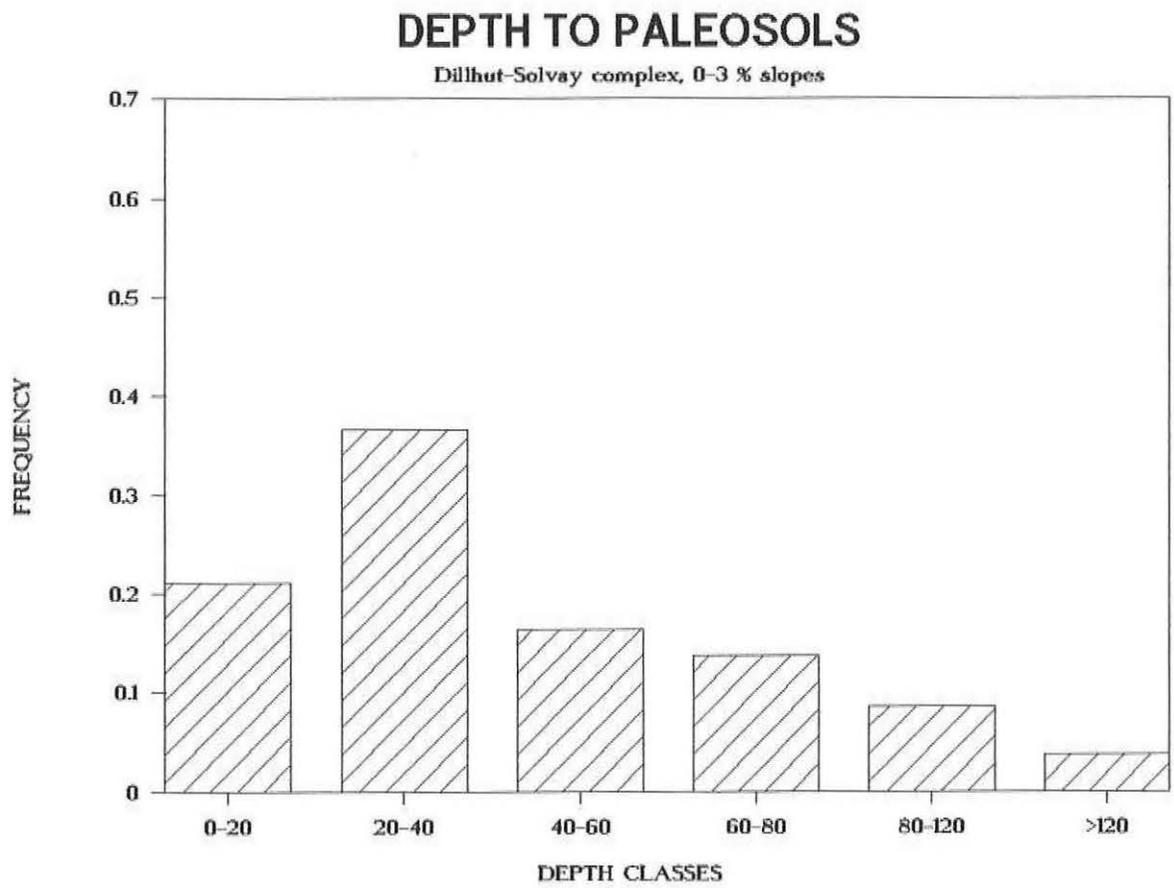
MU: PX

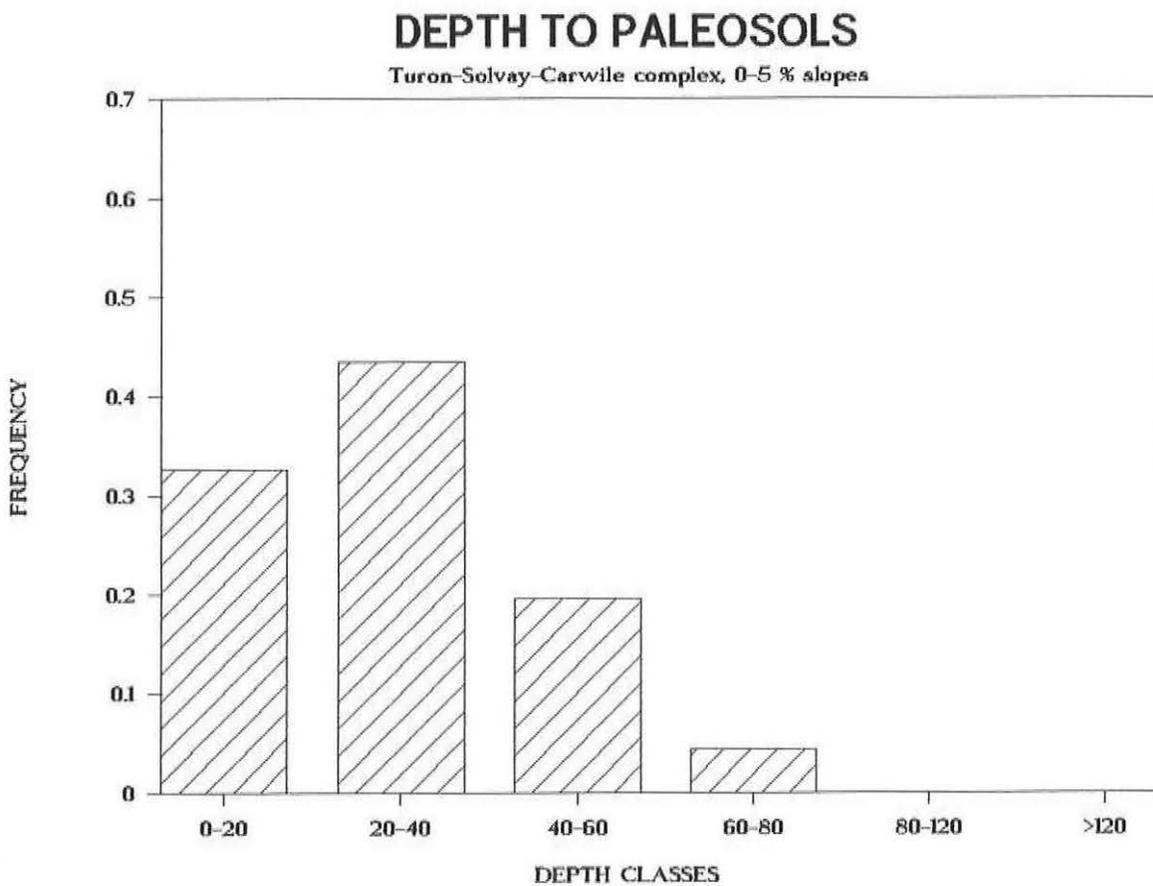
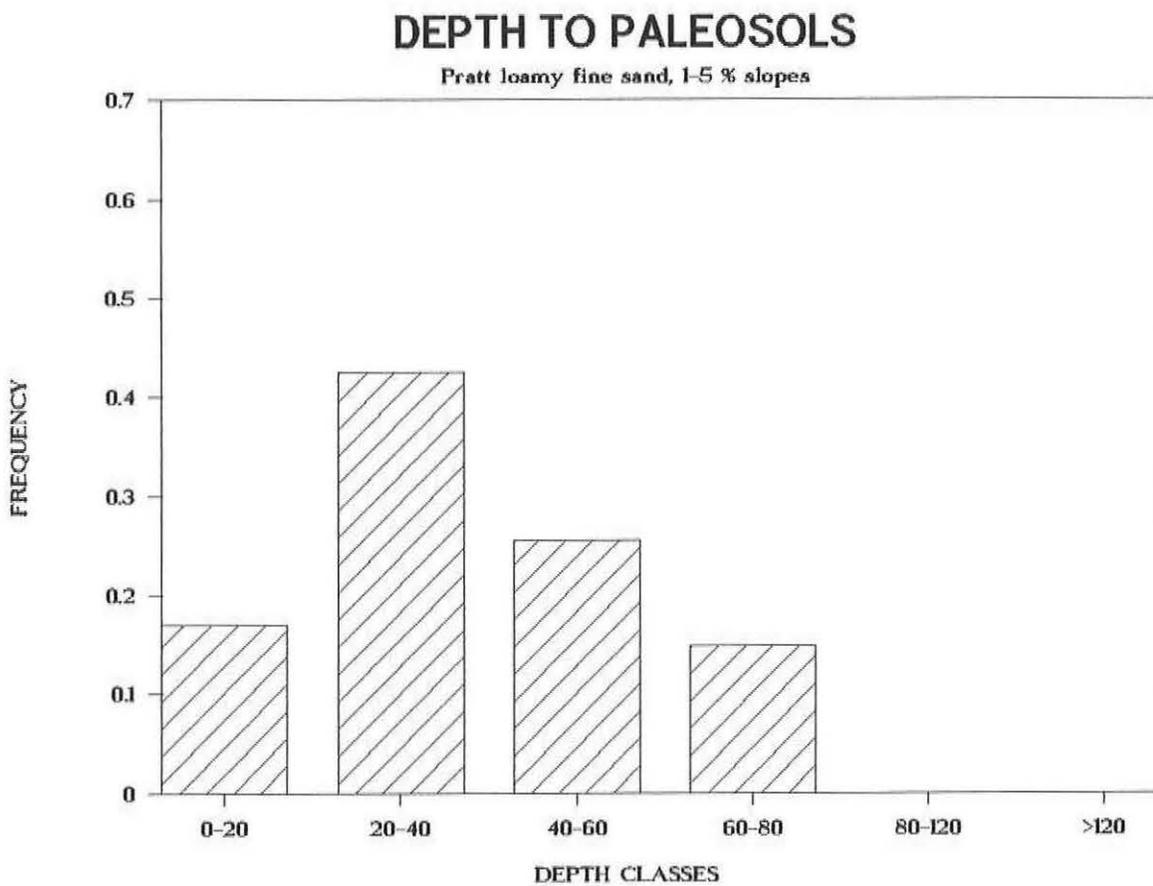
<u>DISTANCE</u>	<u>DEPTH</u>
0'	17"
100	54
200	150
300	94
400	42
500	17
600	49
700	17
800	17
900	17
1000	68
1100	17
1200	19

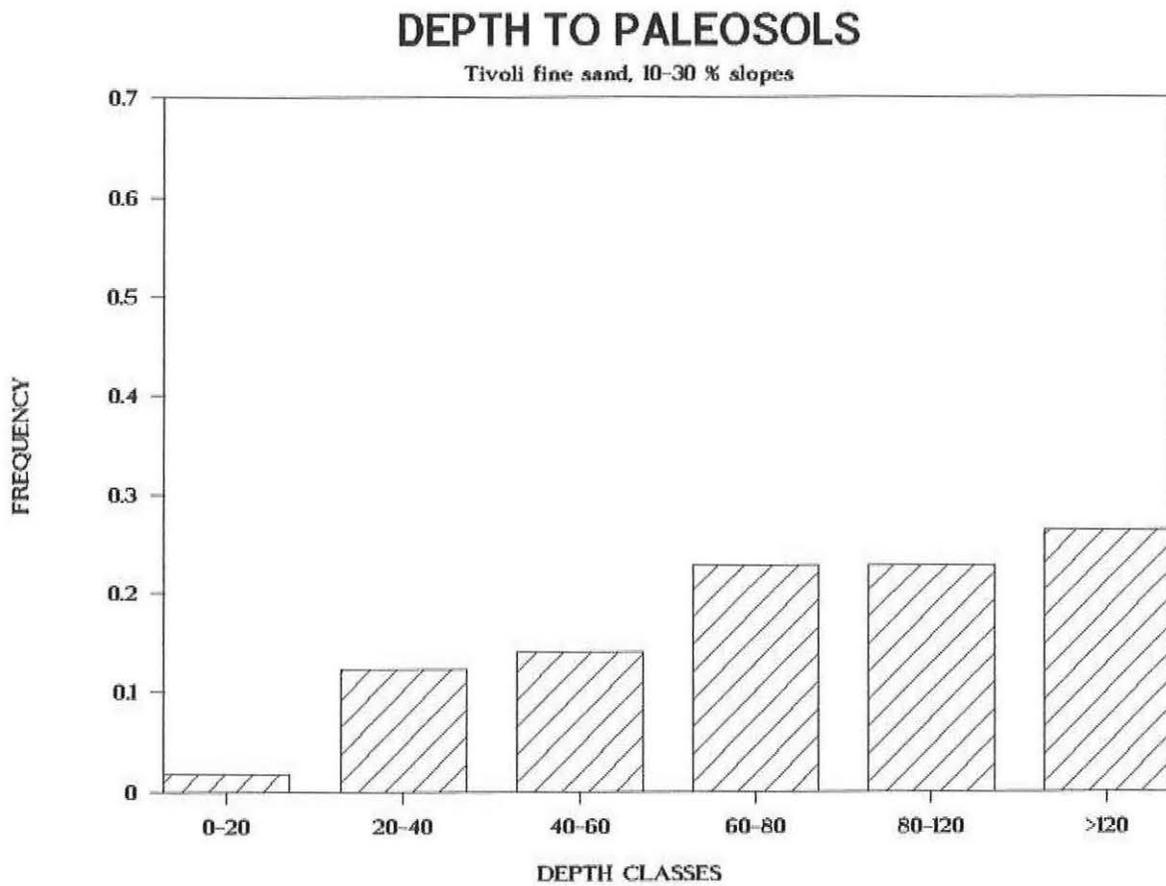
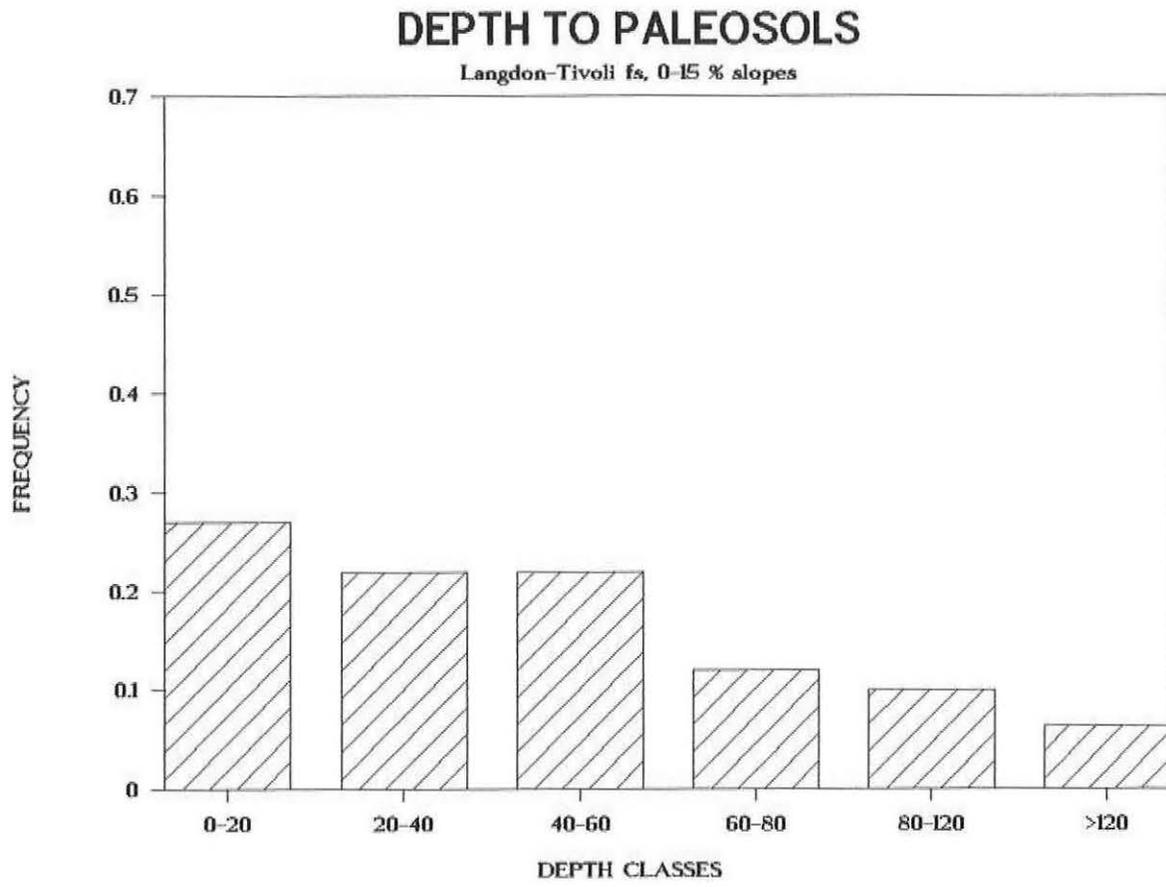
FIGURE 1



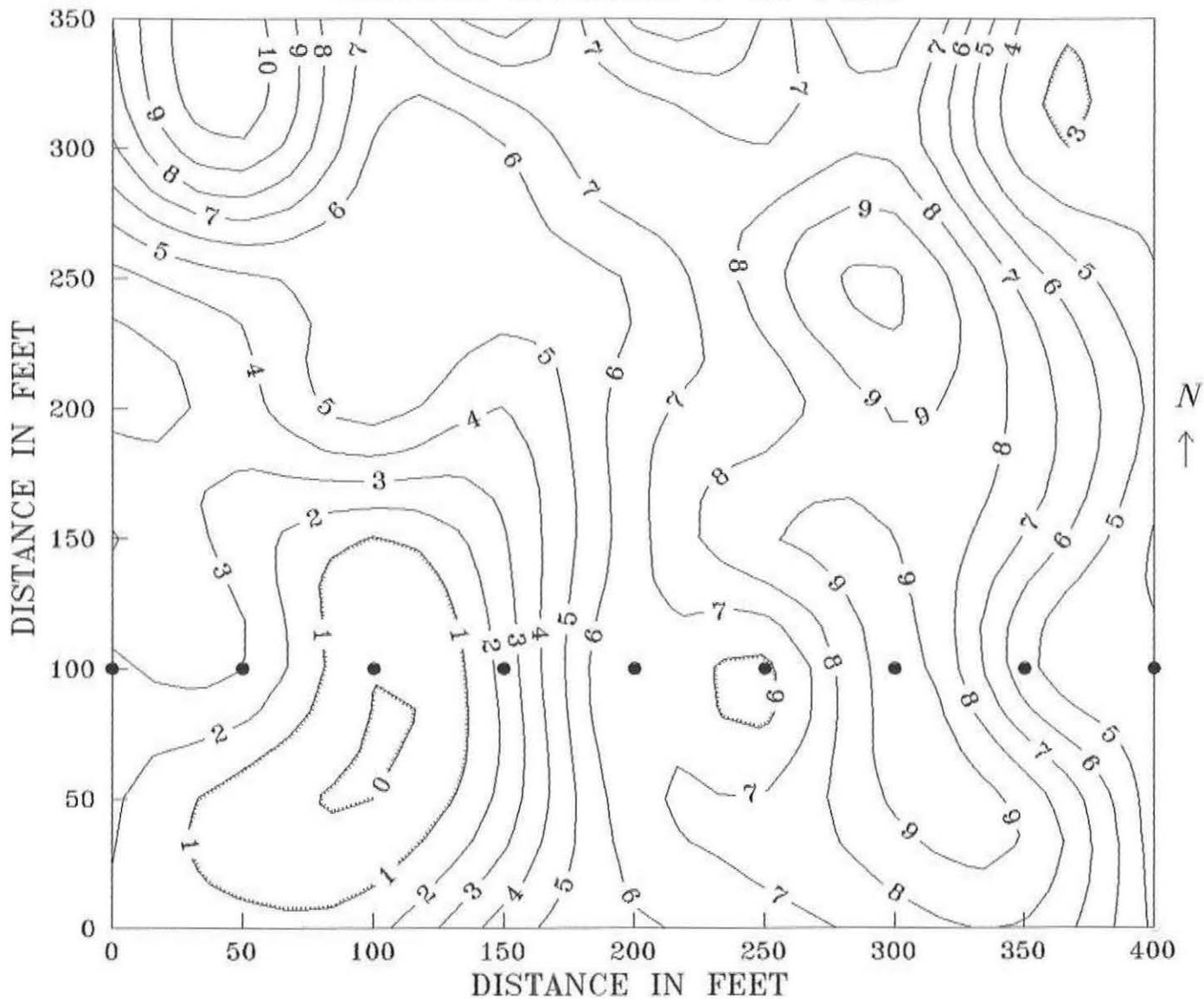




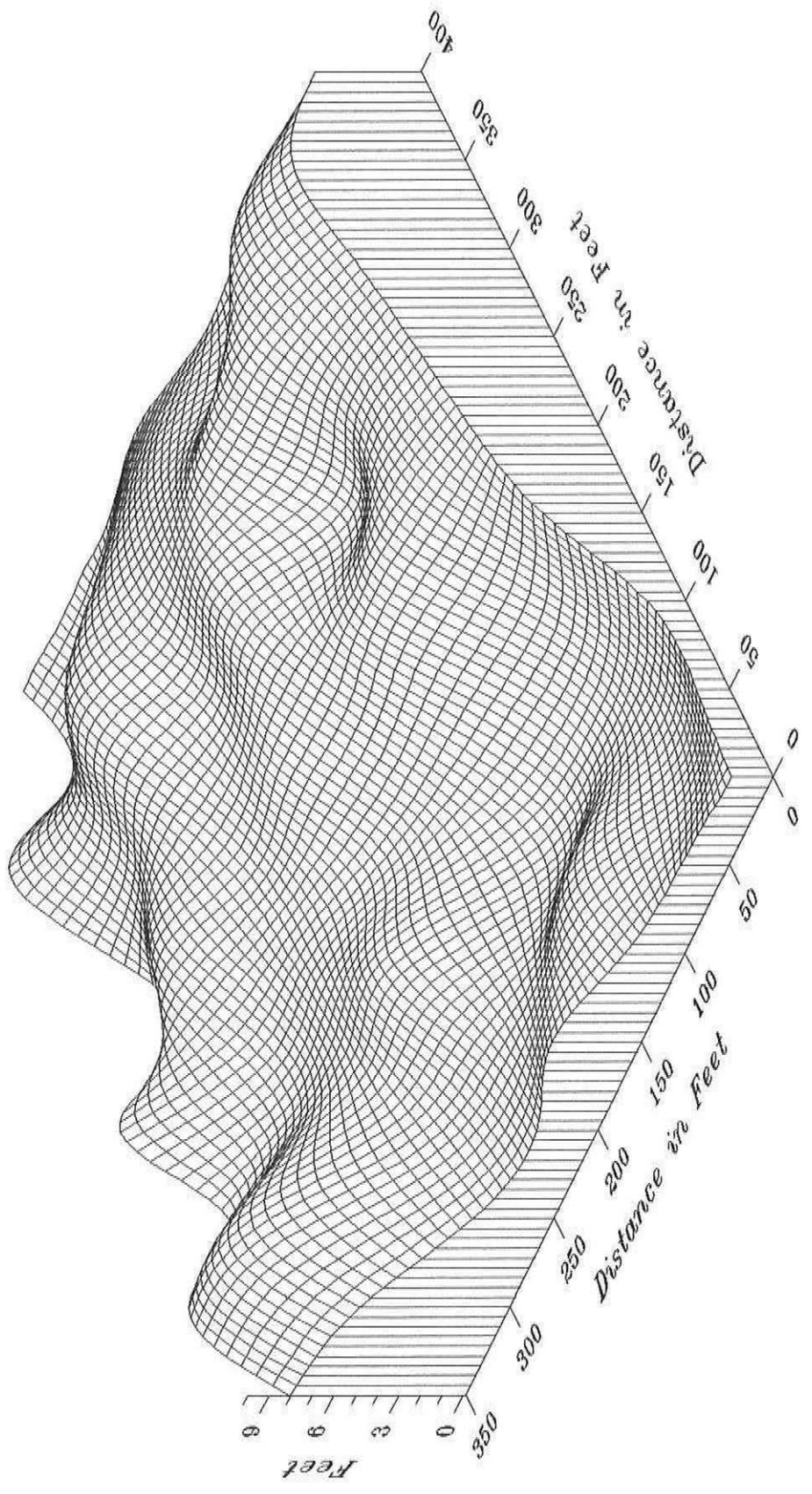




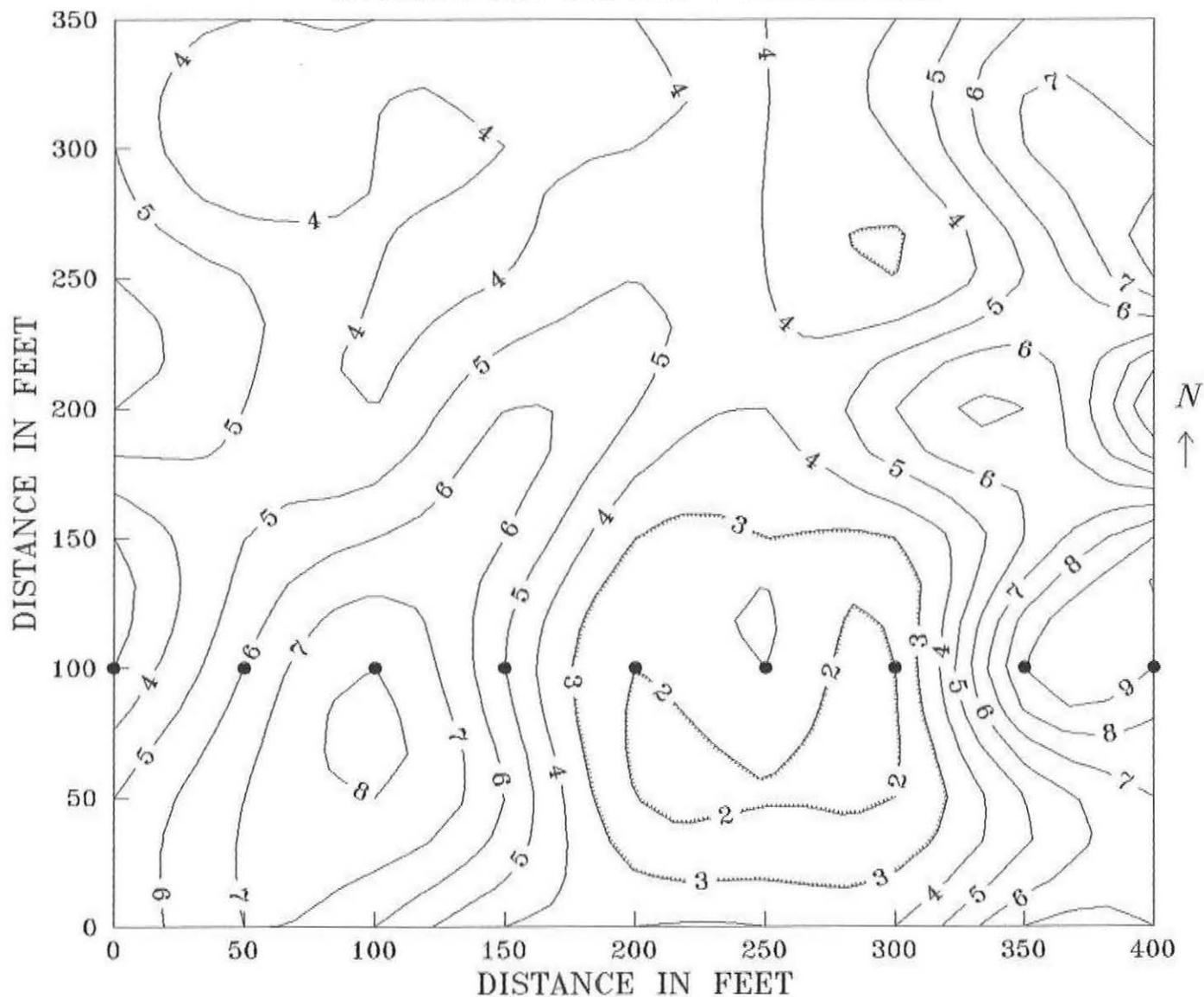
AREA OF LANGDON-TIVOLI FS, 0-15 PERCENT SLOPES  
 TOPOGRAPHY  
 CONTOUR INTERVAL = 1.0 FOOT



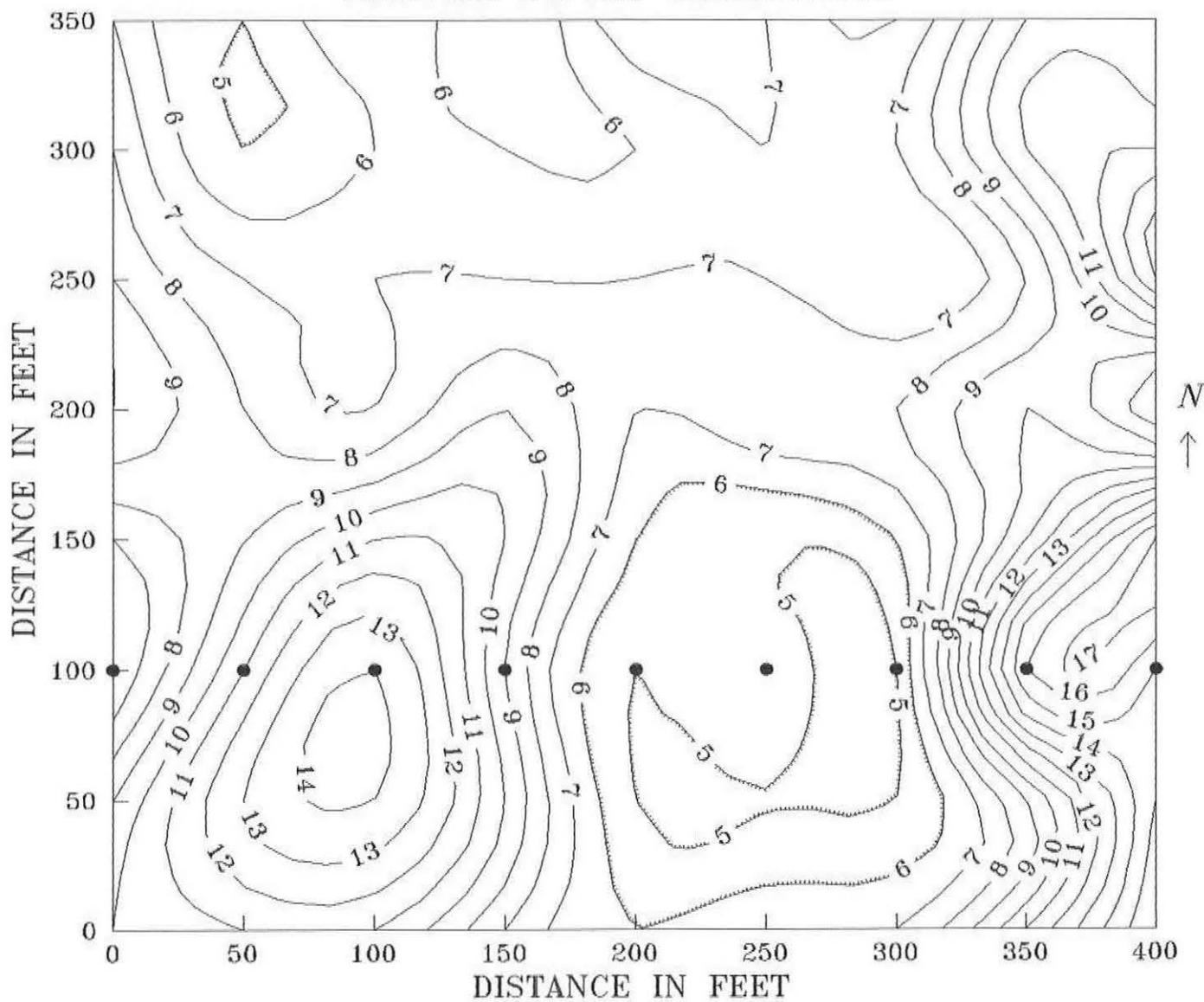
AREA OF LANGDON-TIVOLI FINE SANDS, 0-15 PERCENT SLOPES



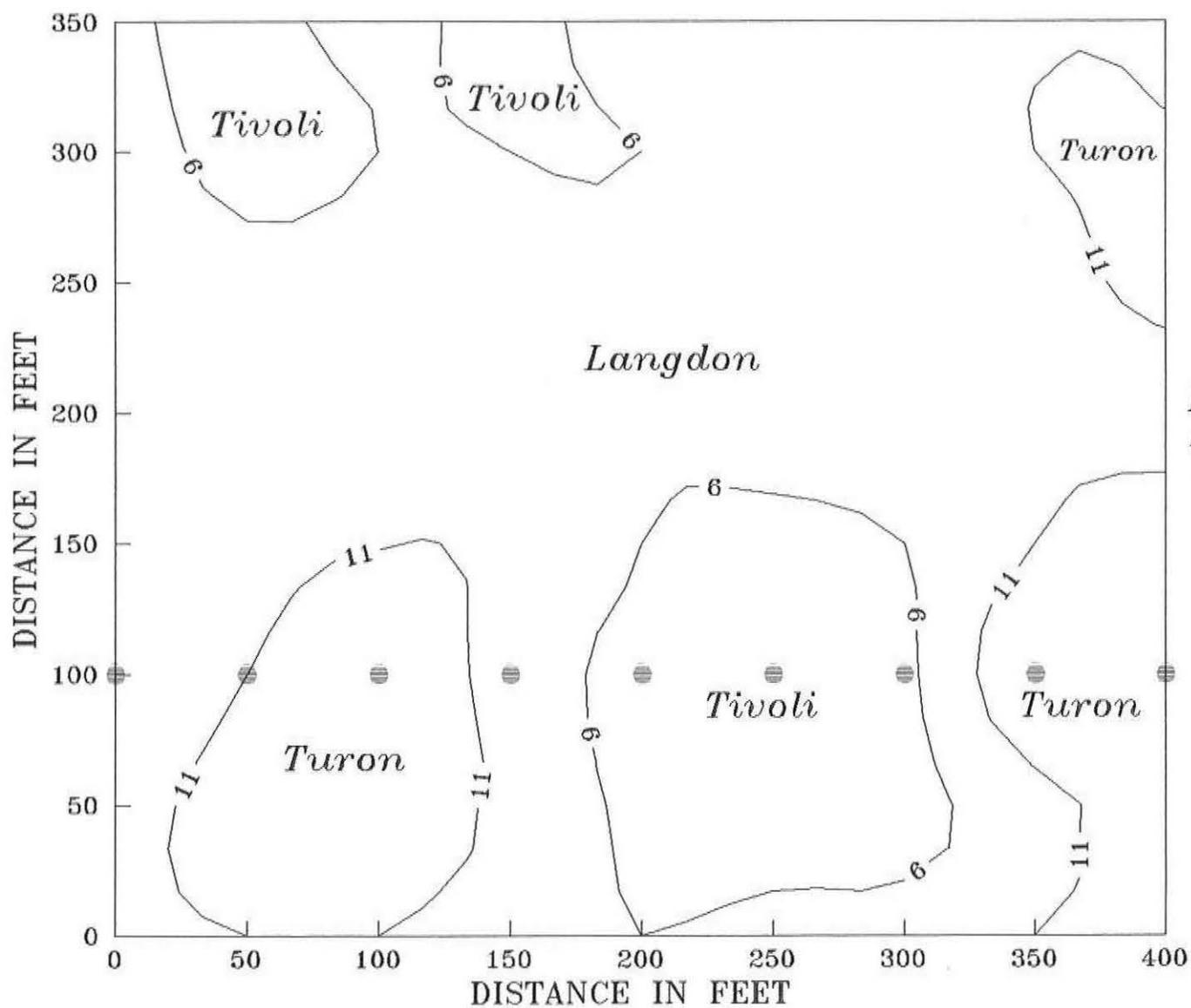
AREA OF LANGDON-TIVOLI FS, 0-15 PERCENT SLOPES  
EM38 SURVEY  
HORIZONTAL DIPOLE ORIENTATION



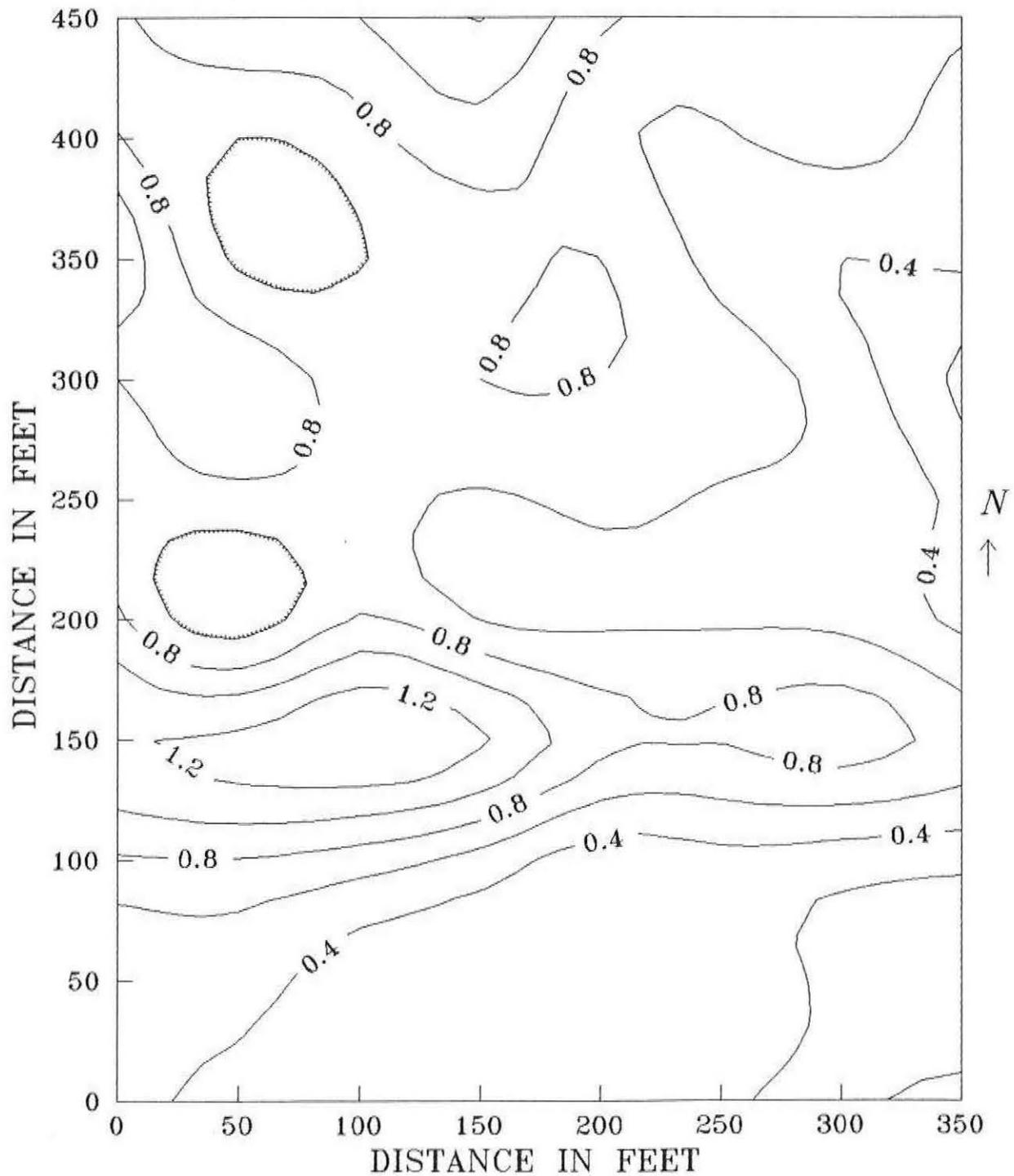
AREA OF LANGDON-TIVOLI FS, 0-15 PERCENT SLOPES  
EM38 SURVEY  
VERTICAL DIPOLE ORIENTATION



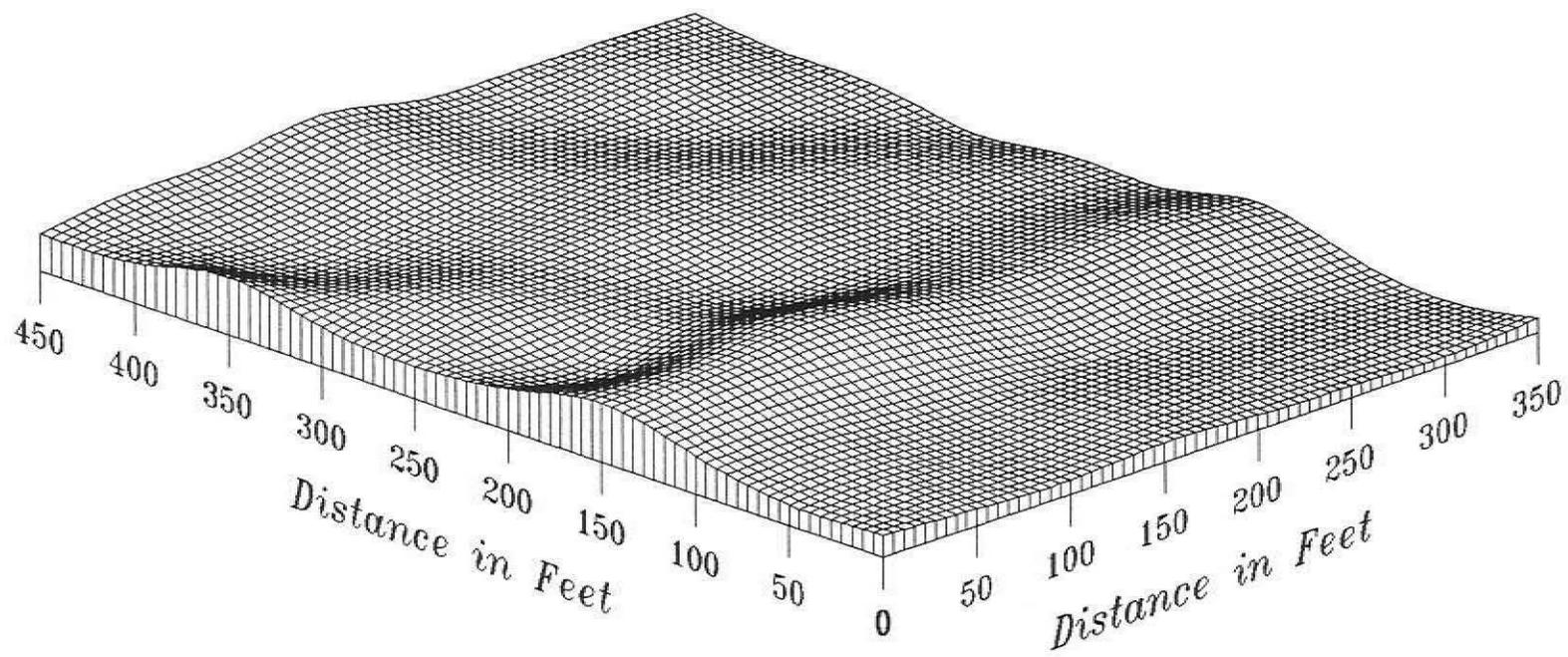
## SOIL MAP PREPARED FROM EM DATA



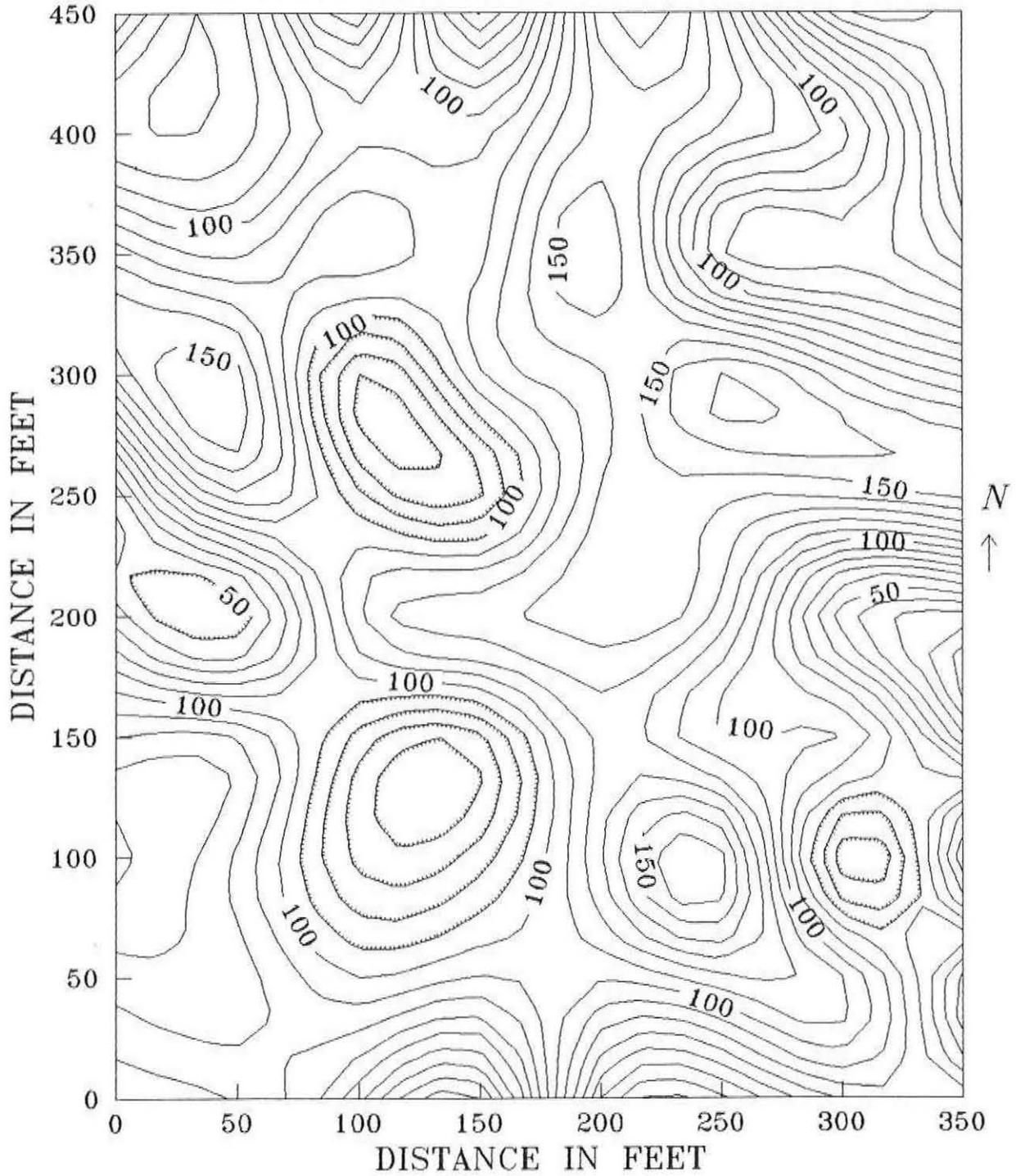
AREA OF TABLER-PUNKIN COMPLEX  
TOPOGRAPHY  
CONTOUR INTERVAL = 0.2 FOOT



## AREA OF TABLER-PUNKIN COMPLEX



AREA OF TABLER-PUNKIN COMPLEX  
EM38 SURVEY  
HORIZONTAL DIPOLE ORIENTATION



AREA OF TABLER-PUNKIN COMPLEX  
*EM38 SURVEY*  
*VERTICAL DIPOLE ORIENTATION*

